

# Modification of the QPE algorithm to account for partial beam blockage

## Update on the HCA and QPE algorithm development

Alexander Ryzhkov, NSSL

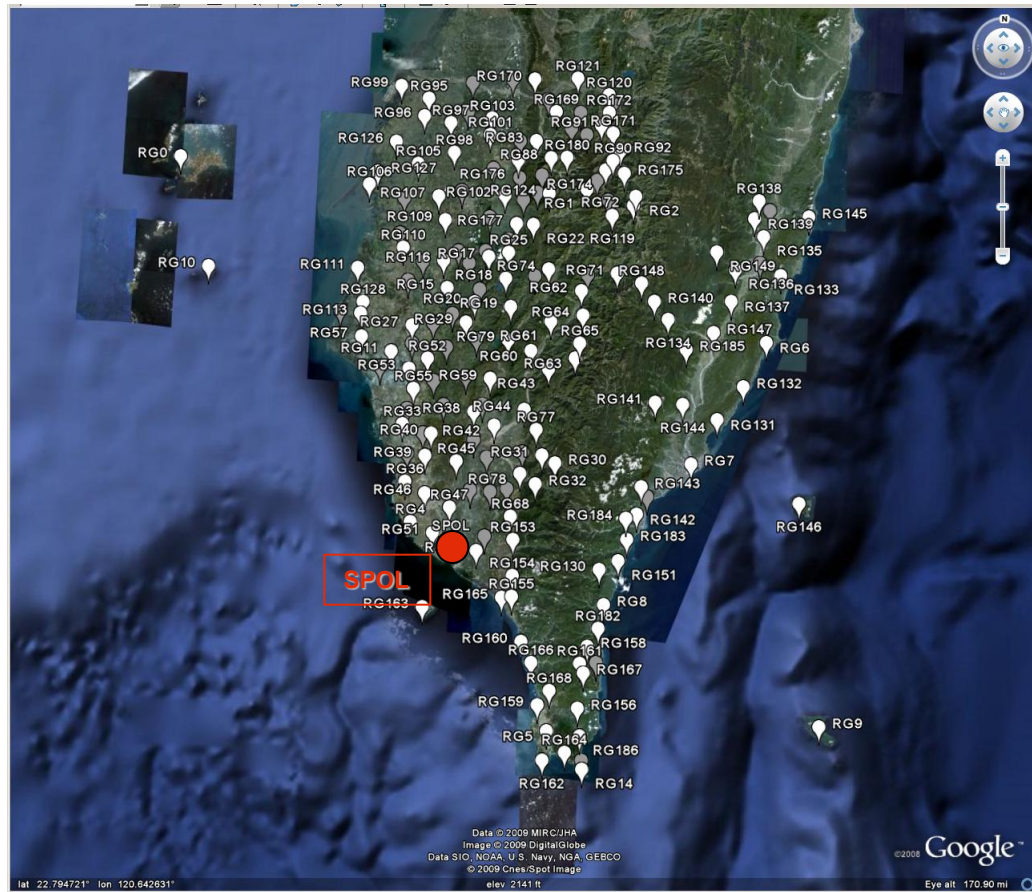
November 19, 2009

# Outline of the talk

1. Investigation of the impact of complex terrain / partial beam blockage on the quality of polarimetric rainfall measurements using SPOL data collected in Taiwan
2. Brief summary of recent research efforts to improve polarimetric HCA and QPE
  - Physical model-based polarimetric VPR
  - Performance of the polarimetric QPE algorithm in the areas affected by ground clutter canceling
  - Modification of HCA to discriminate between small and large hail
  - Developing HCA module for transitional winter weather

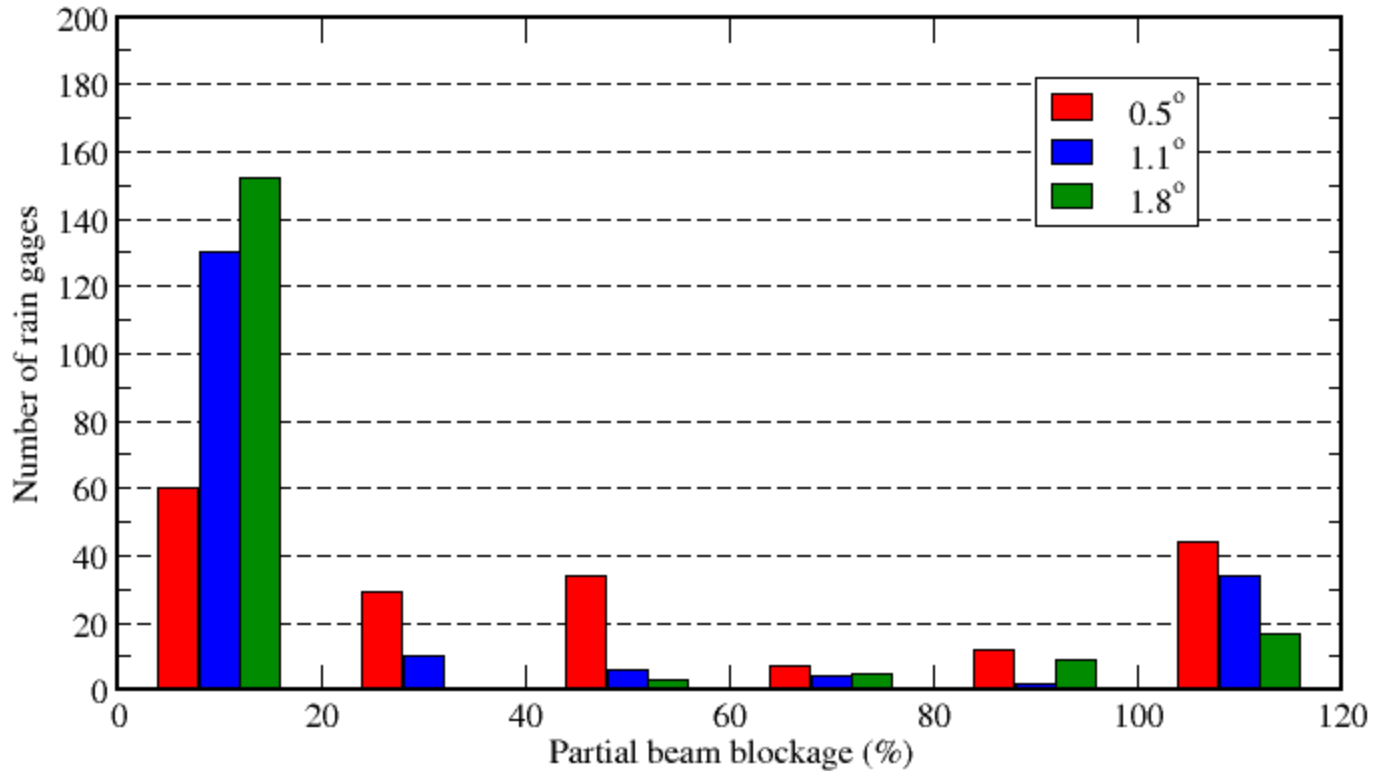
# SoWMEX / TiMREX experiment in Taiwan, 2008.

## Locations of Taiwan rain gages and SPOL radar

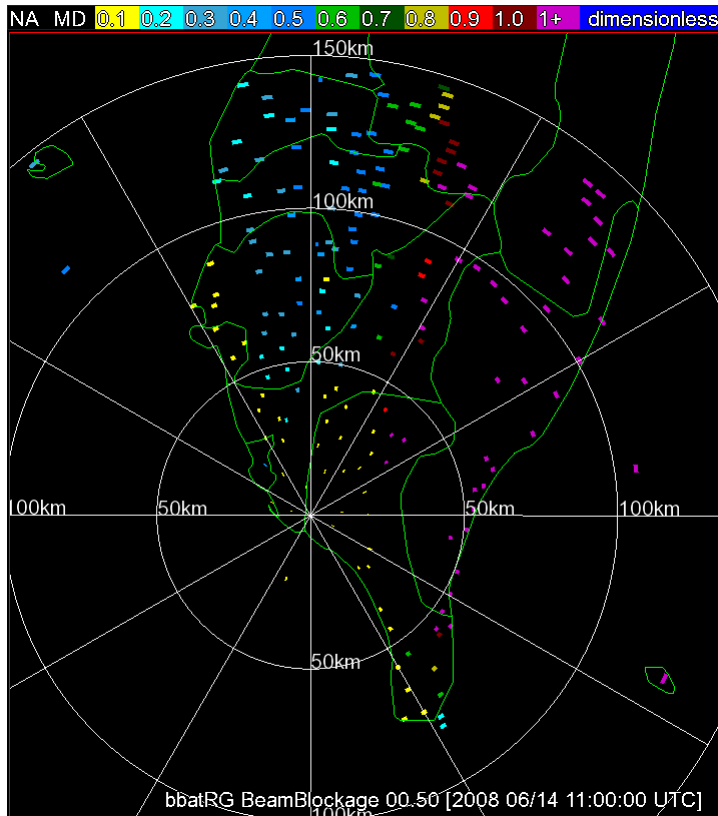


$R < 150 \text{ km}$

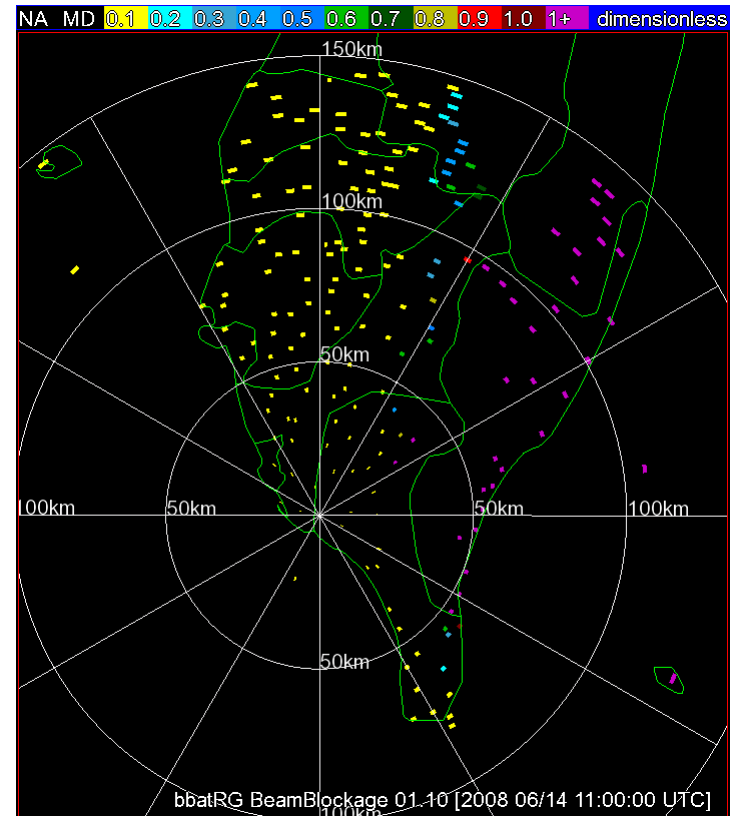
# Degree of blockage at different elevations for raingage sites in Taiwan with respect to the SPOL radar



# Partial Beam Blockage at Different Elevation Angles

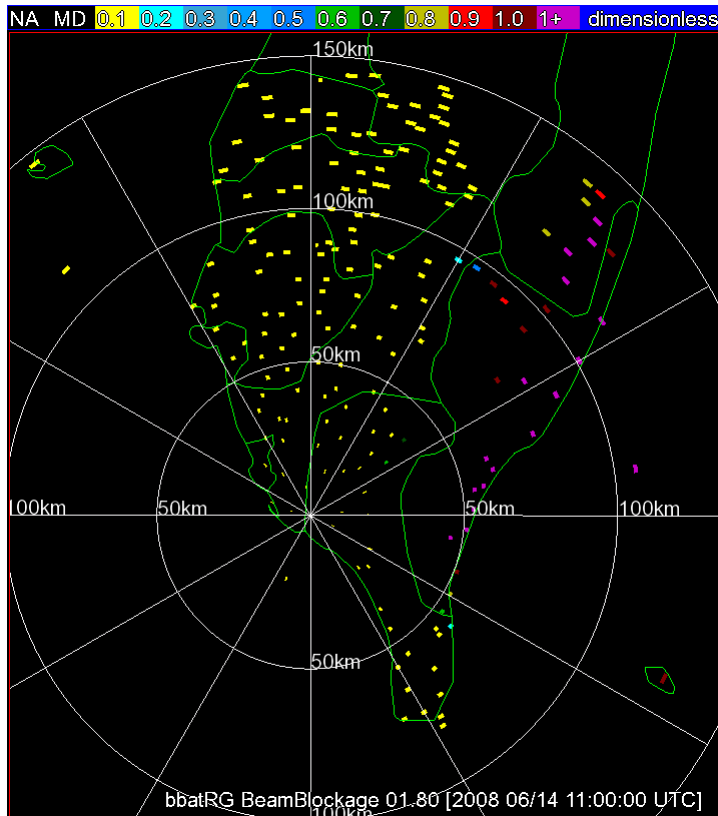


0.5

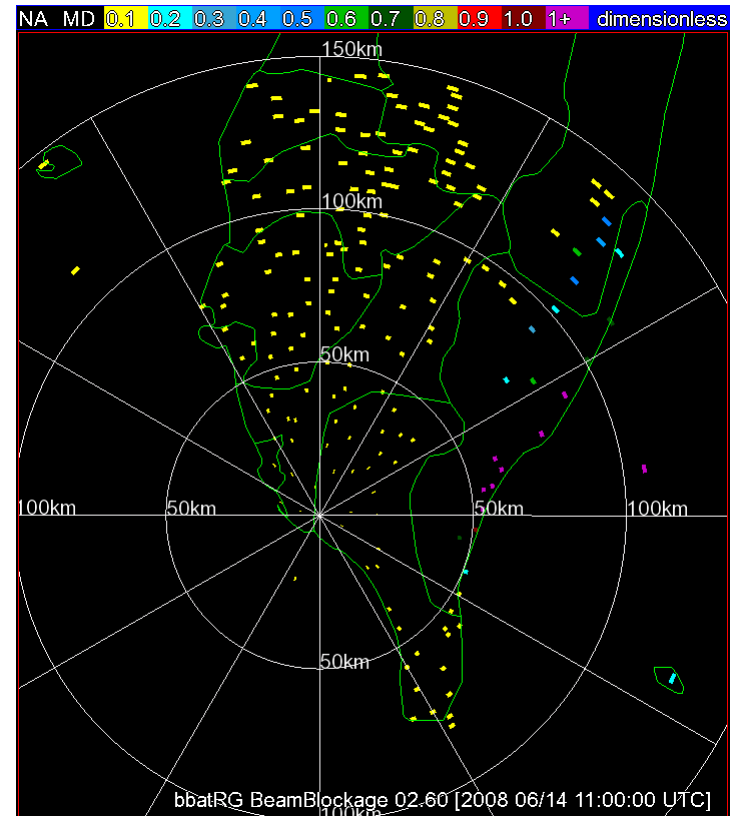


1.1

# Partial Beam Blockage at Different Elevation Angles



1.8

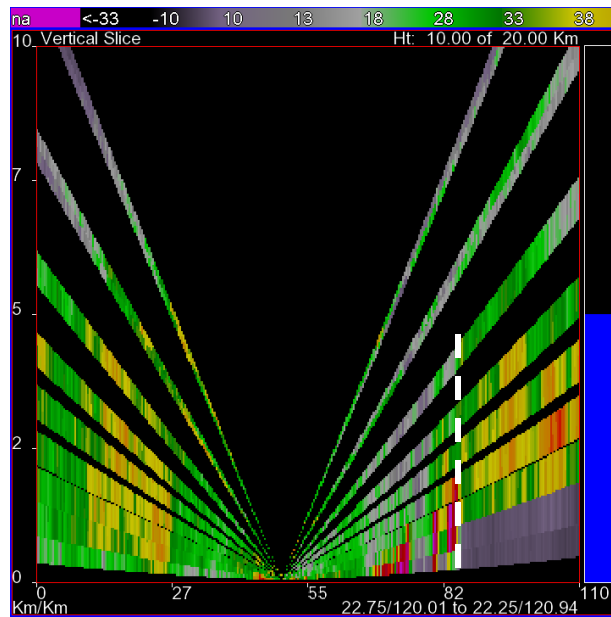


2.6

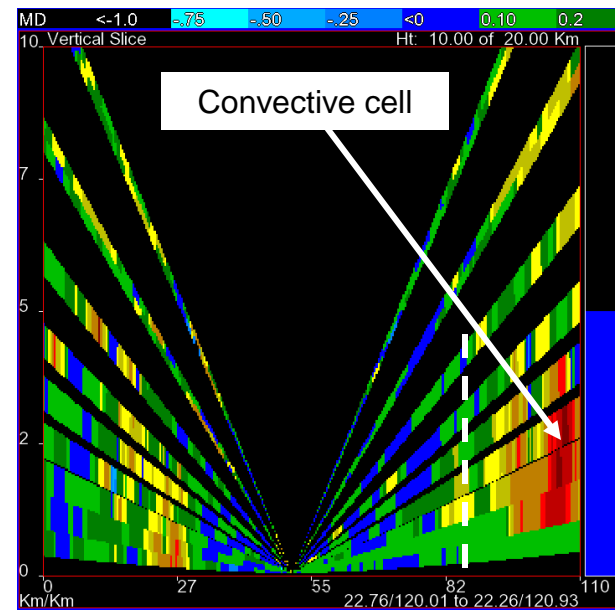
# How can $K_{DP}$ help?

RHI plots of  $Z$  and  $K_{DP}$  measured by the NCAR SPOL radar in Taiwan

Radar reflectivity  $Z$



Specific differential phase  $K_{DP}$



$K_{DP}$  "senses" convective cell behind the mountain better than  $Z$

# List of rain events

1. 2008/06/02 21 – 22 UTC
2. 2008/06/05 01 – 03 UTC
3. 2008/06/14 10 – 13 UTC



# Four different algorithms for rainfall estimation

$$R_1(Z) = 1.7010^{-2} 10^{0.0714Z} \quad \text{Standard WSR-88D}$$

$$R_2(Z) = R_1(Z) F^{-0.714} \quad \text{Standard WSR-88D with geometrical blockage correction}$$

$$F = 0.5 \tanh[0.0277(50 - b)] + 0.5$$

$$R_3(Z, Z_{DR}) = 1.33 R_2^{1.30}(Z) 10^{-0.343Z_{DR}} \quad \text{Bringi and Chandra, 2001}$$

$$R_4(Z, Z_{DR}, K_{DP}) = \frac{R_2(Z)}{0.4 + 5.0 |Z_{dr} - 1|^{1.3}} \quad \text{if } R_1(Z) < 6 \text{ mm / h}$$

$$R_4(Z, Z_{DR}, K_{DP}) = \frac{R(K_{DP})}{0.4 + 3.5 |Z_{dr} - 1|^{1.7}} \quad \text{if } R_1(Z) > 6 \text{ mm / h}$$

$$R(K_{DP}) = 44.0 |K_{DP}|^{0.822} \text{sign}(K_{DP}) \quad \text{Synthetic (Ryzhkov et al. 2005)}$$

# Correction of Z based on the geometry of obstruction

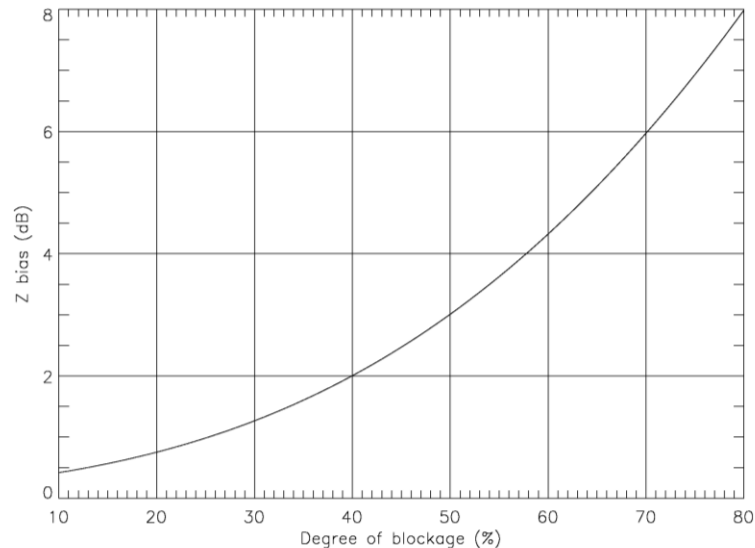
$$\alpha = 100 \frac{\theta_b - \theta_0 + \Omega/2}{\Omega}$$

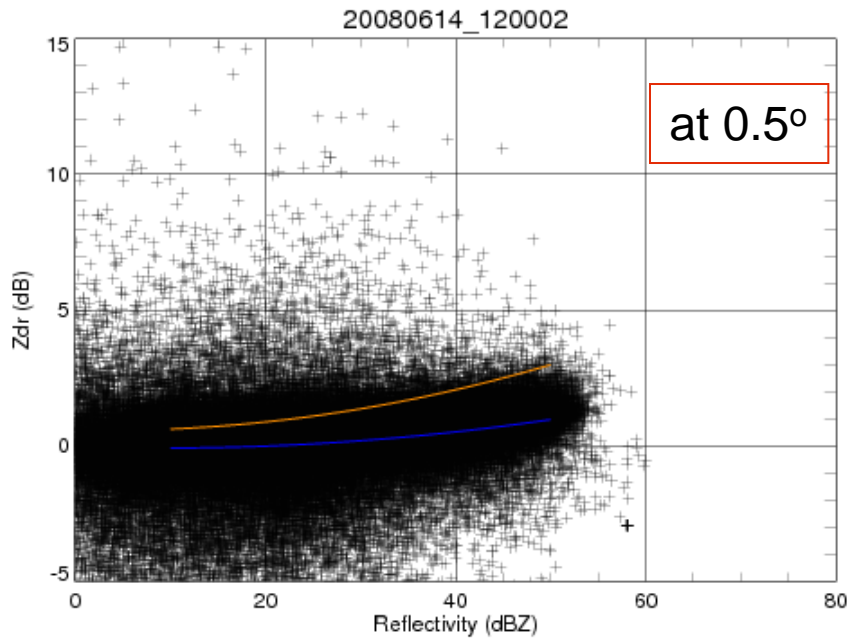
$\alpha$  is blockage degree (%)

$\Omega$  is radar beamwidth,  $\theta_0$  is elevation of the beam center,  $\theta_b$  is blockage elevation

$$Z(\text{dBZ}) = Z_{\text{blocked}}(\text{dBZ}) + 10 \log(F_{\text{shield}})$$

$$F_{\text{shield}} = 0.5 \tanh[0.0277(50 - \alpha)] + 0.5 \quad \text{if } \Omega = 1.0$$

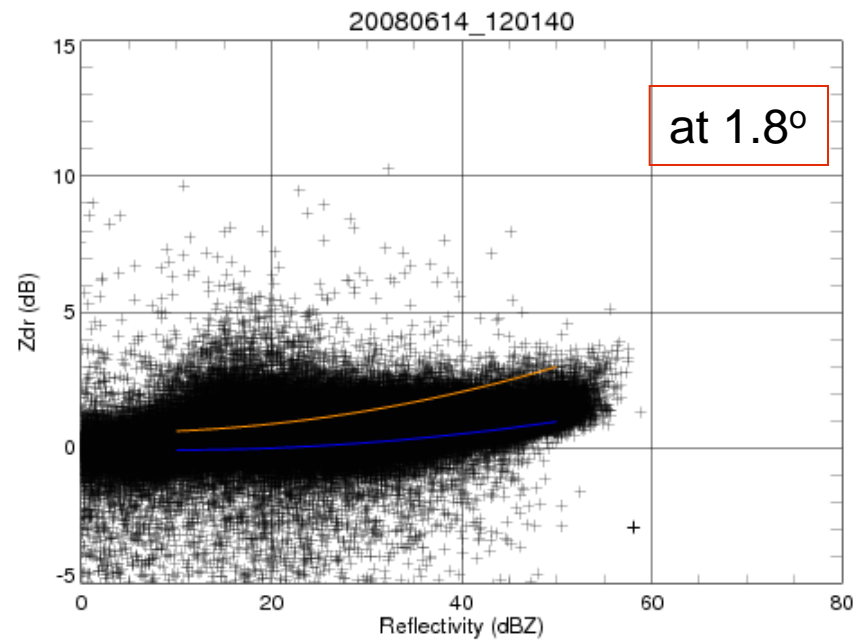
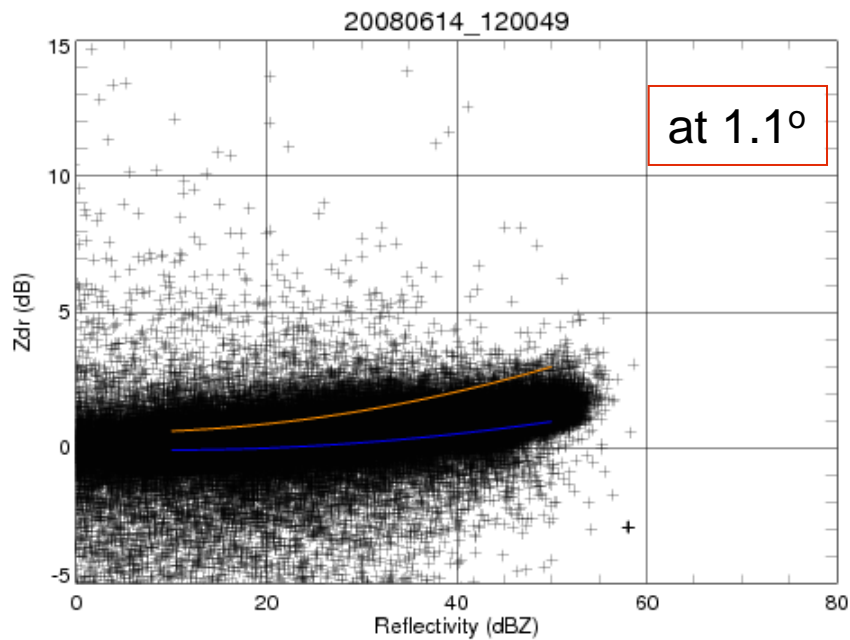




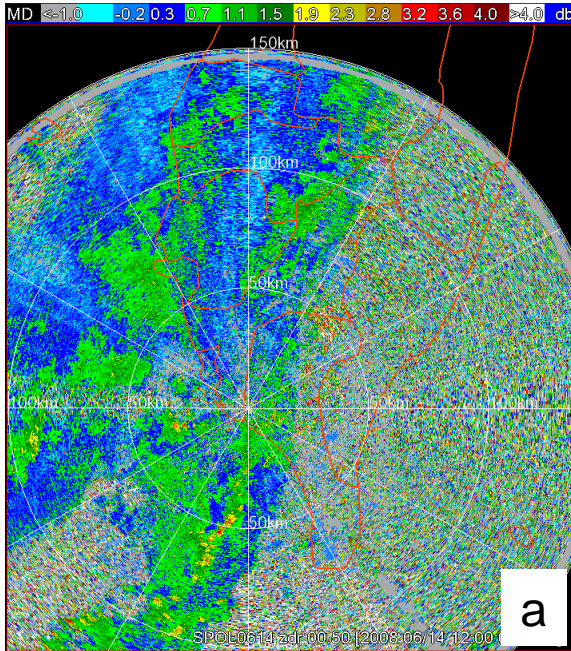
Z – Z<sub>DR</sub> scatterplots at elevations  $0.5^\circ$ ,  $1.1^\circ$ , and  $1.8^\circ$

**Z<sub>DR</sub> is negatively biased at lowest elevation due to contamination from ground clutter**

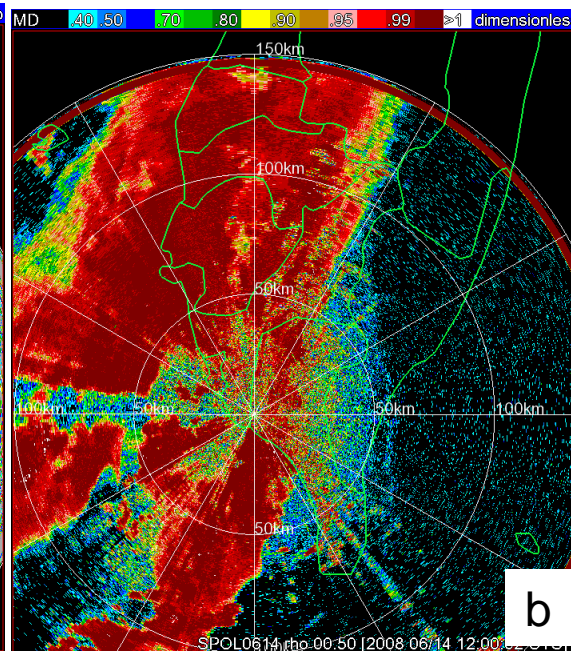
With  $\rho_{hv} > 0.8$ ,  $r < 140$  km



$Z_{dr}$

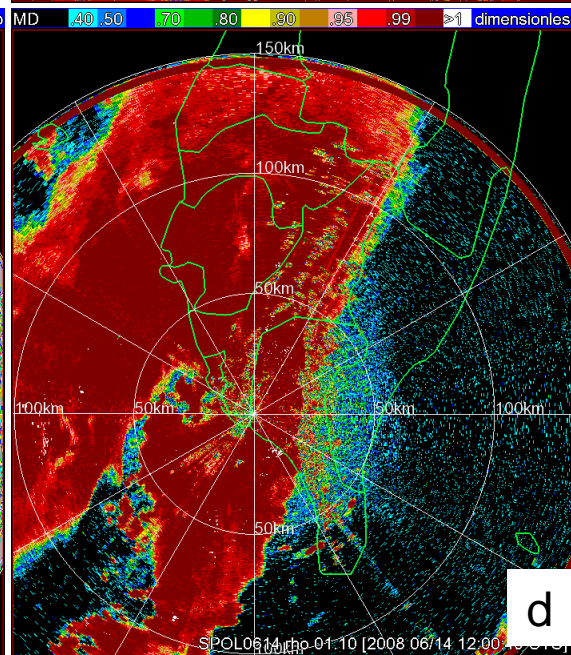
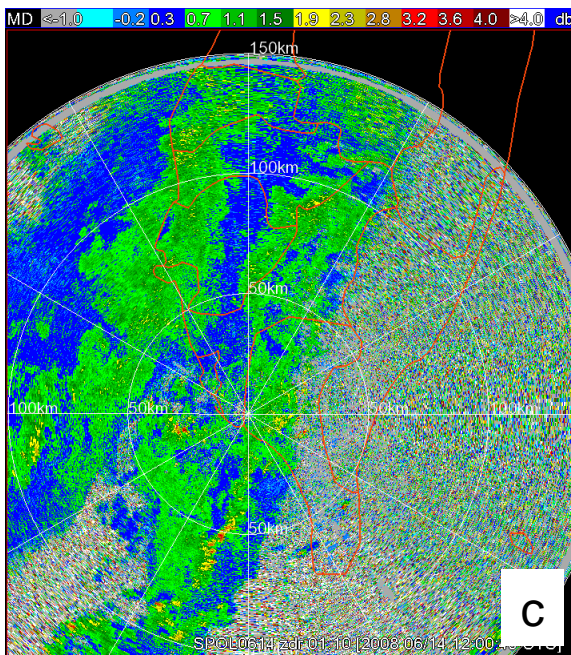


$\rho_{hv}$



$Z_{DR}$  and  $\rho_{hv}$  are lower at lower elevation

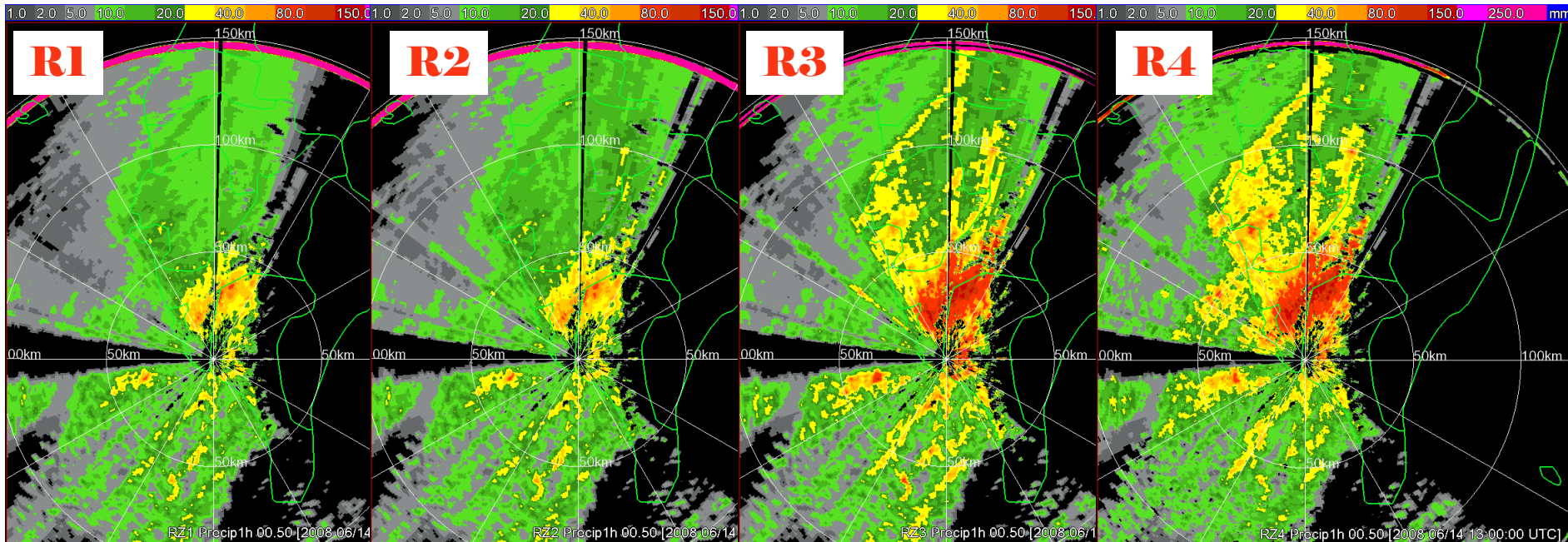
At  $0.5^\circ$



At  $1.1^\circ$

# 3 hour rain totals computed using 4 algorithms at elevation 0.5

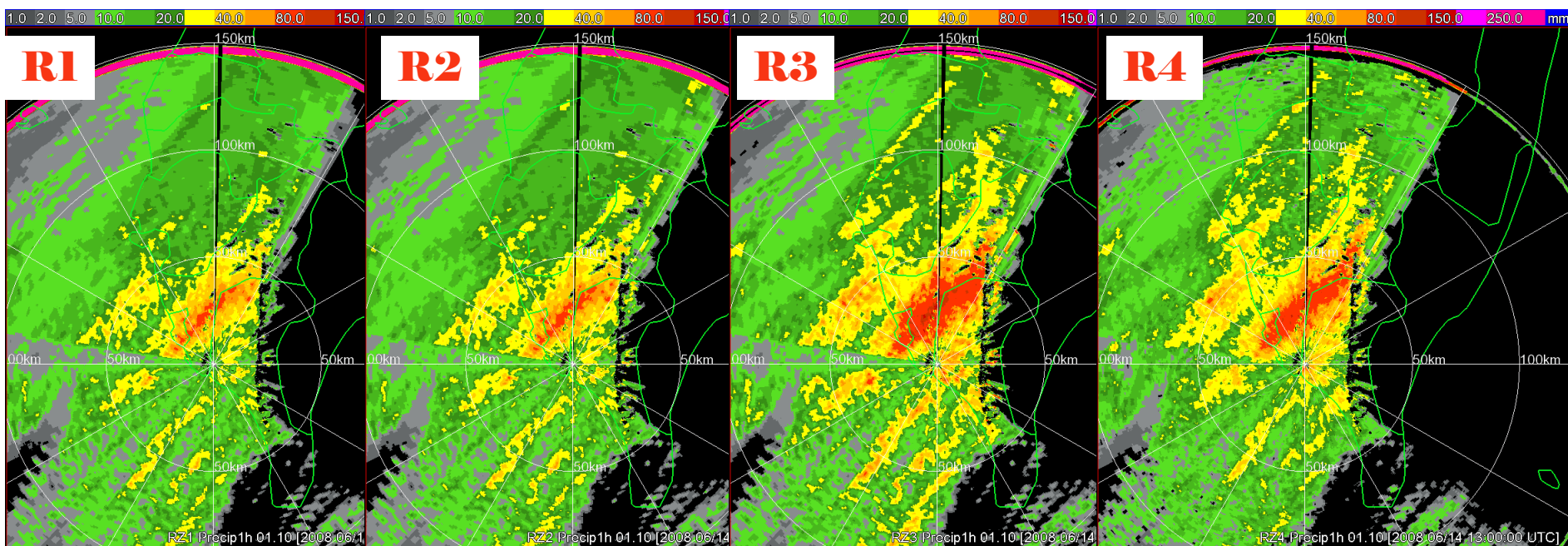
2008/06/14, 10 – 13 UTC



$R_1$  and  $R_2$  – conventional relations,  $R_3$  and  $R_4$  – polarimetric relations

# 3 hour rain totals computed using 4 algorithms at elevation 1.1

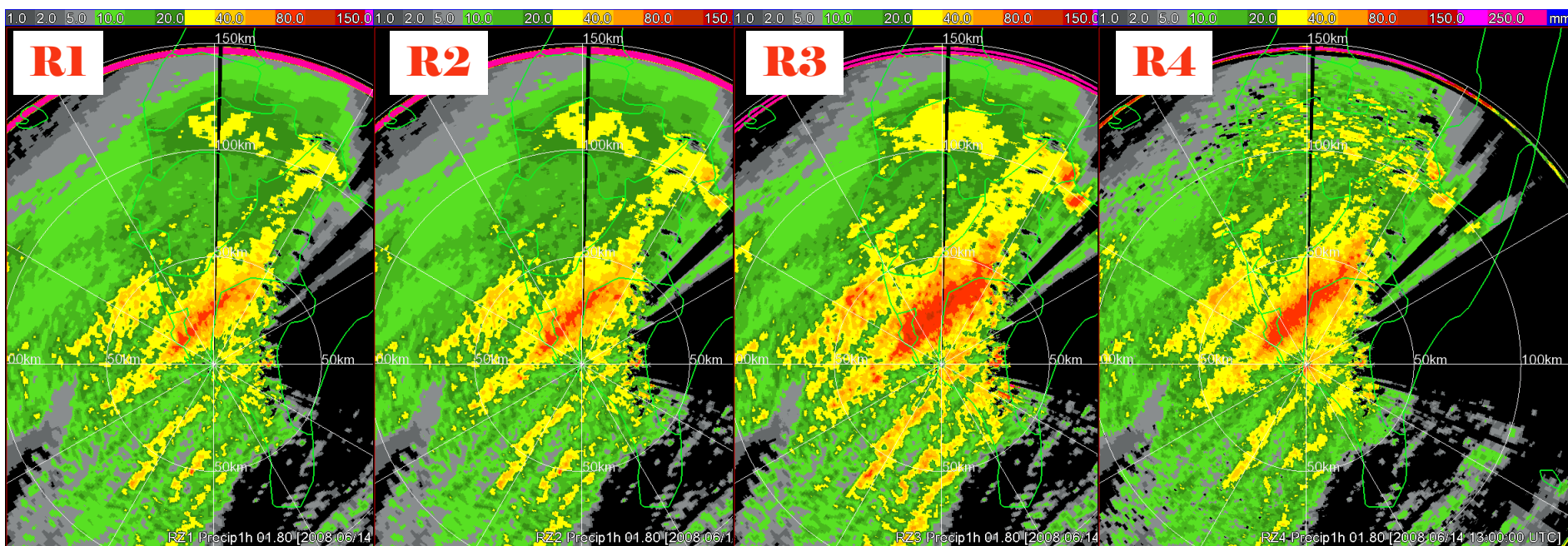
2008/06/14, 10 – 13 UTC



$R_1$  and  $R_2$  – conventional relations,  $R_3$  and  $R_4$  – polarimetric relations

# 3 hour rain totals computed using 4 algorithms at elevation 1.8

2008/06/14, 10 – 13 UTC



$R_1$  and  $R_2$  – conventional relations,  $R_3$  and  $R_4$  – polarimetric relations

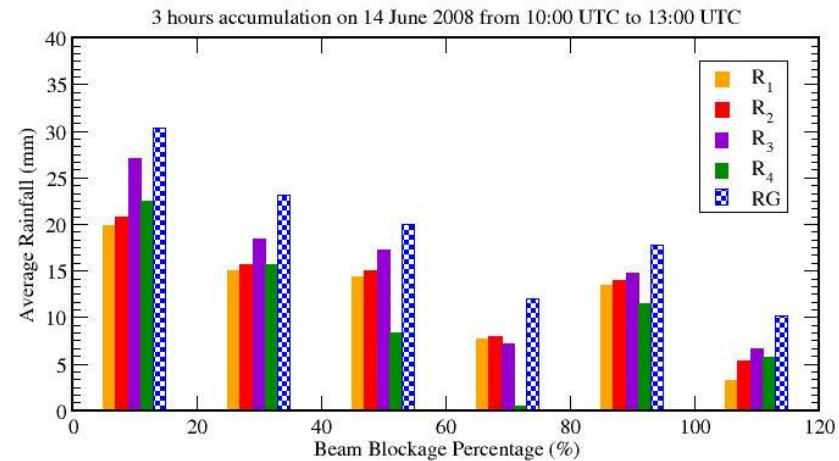
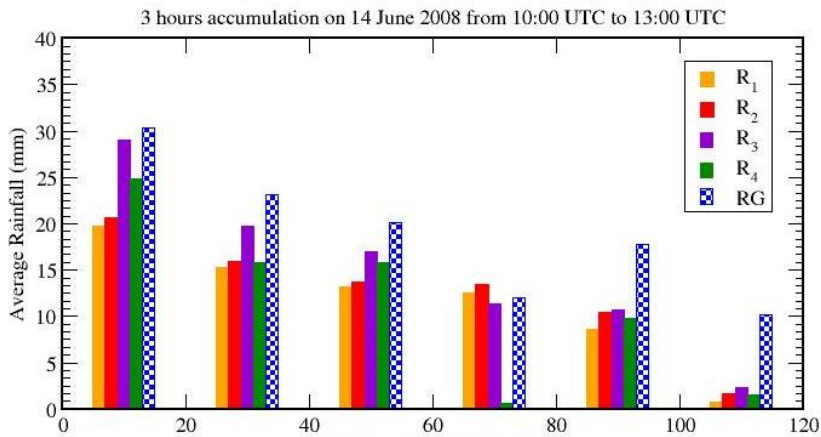
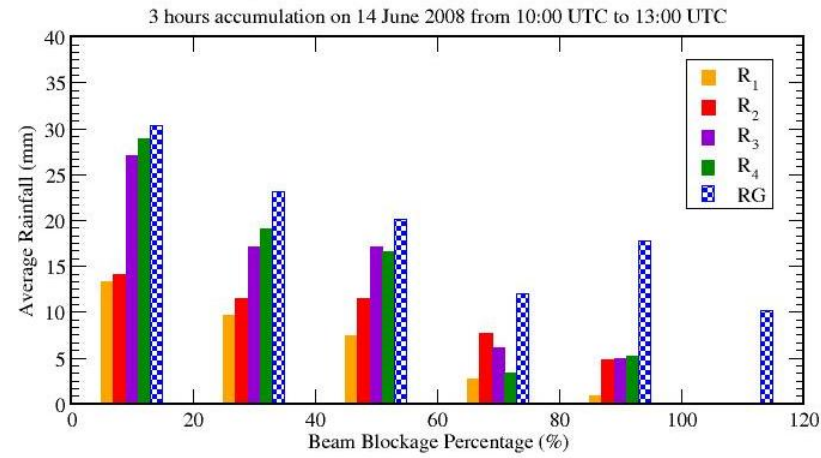
Average rain 3 hr total (mm)

2008/06/14 10 – 13 UTC

EI = 0.5

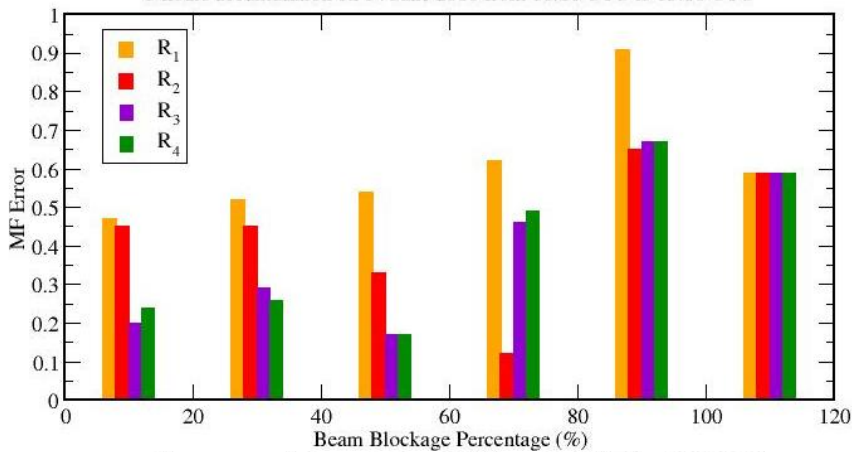
EI = 1.1

EI = 1.8





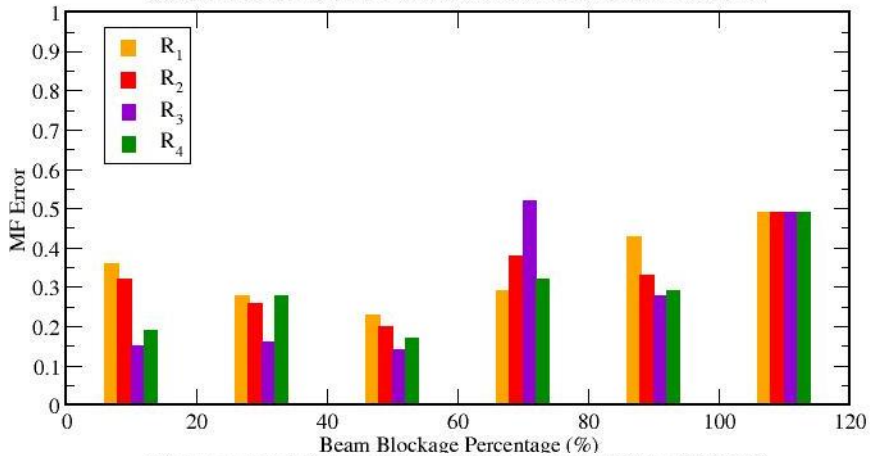
3 hours accumulation on 14 June 2008 from 10:00 UTC to 13:00 UTC



EI = 0.5

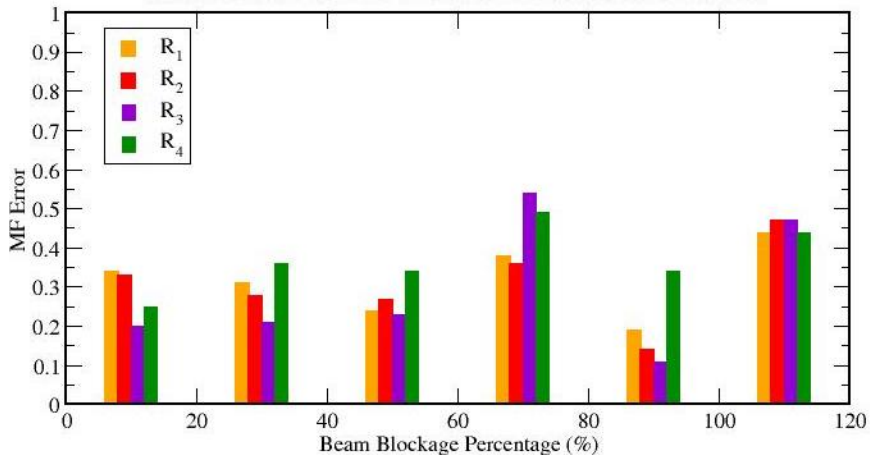
Median fractional error  
2008/06/14 10 – 13 UTC

3 hours accumulation on 14 June 2008 from 10:00 UTC to 13:00 UTC



EI = 1.1

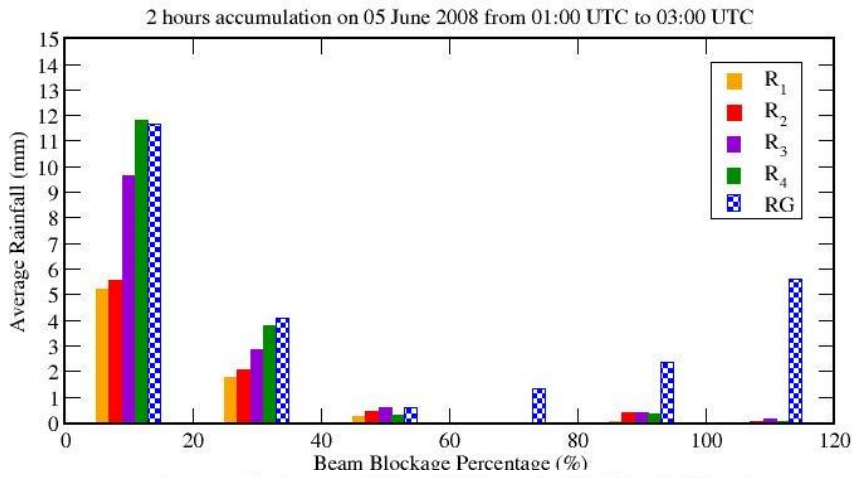
3 hours accumulation on 14 June 2008 from 10:00 UTC to 13:00 UTC



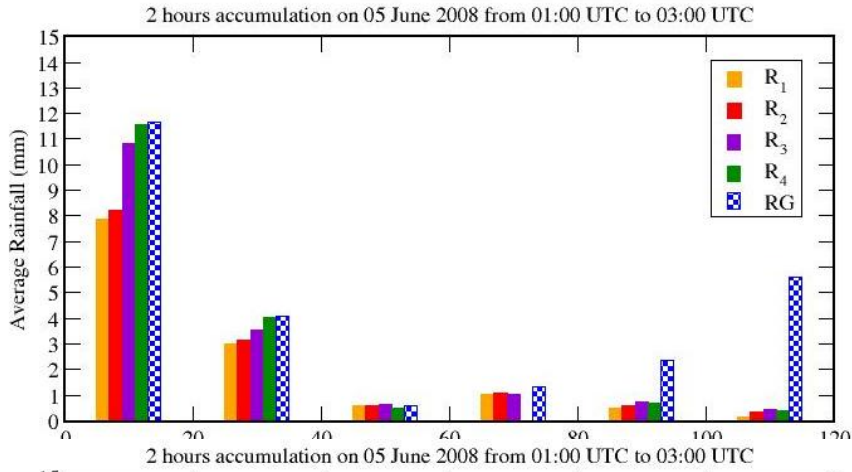
EI = 1.8

Average rain 3 hr total (mm)

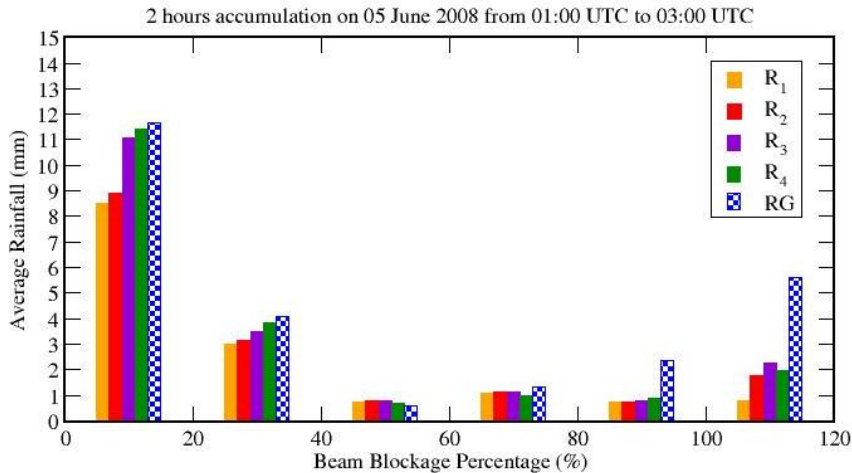
2008/06/05 1 – 3 UTC



EI = 0.5

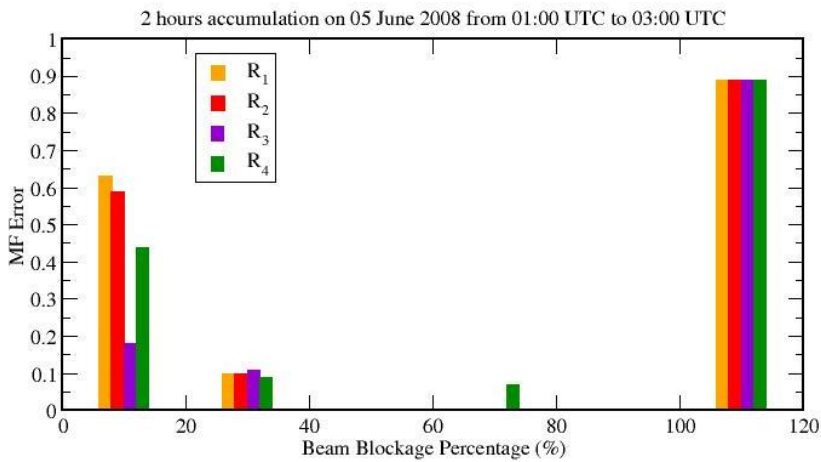


EI = 1.1

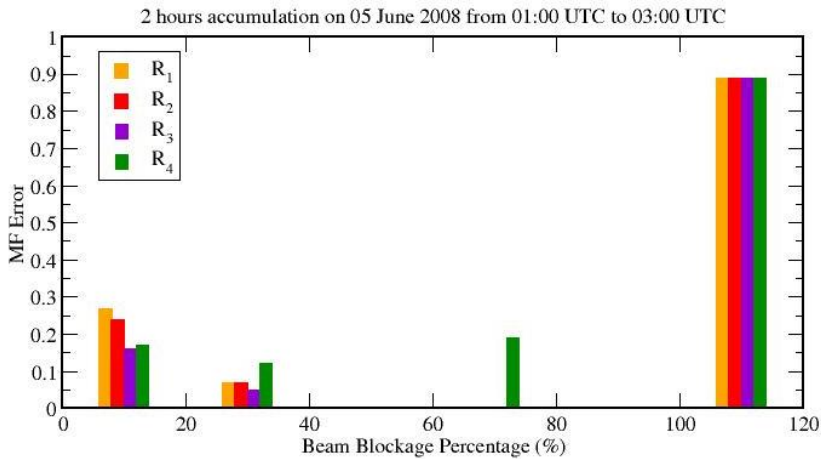


EI = 1.8

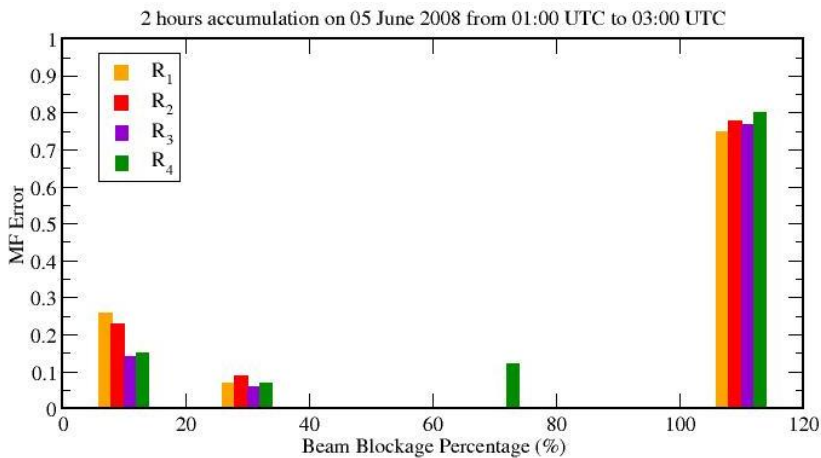
# Median fractional error 2008/06/05 1 – 3 UTC



EI = 0.5



EI = 1.1



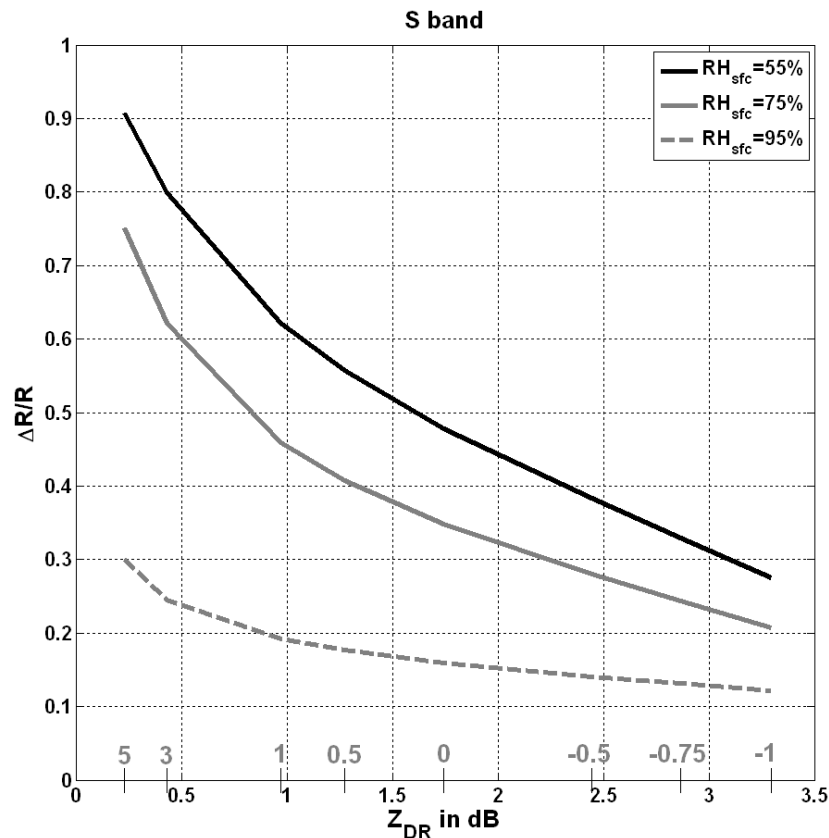
EI = 1.8

# Retrieval of vertical profile of rain rate using microphysical models

## Impact of evaporation

$R < 5 \text{ mm/h}$

Relative change in rain rate due to evaporation as a function of  $Z_{DR}$  aloft and surface relative humidity



Input parameters:

$Z$  and  $Z_{DR}$  at lowest unobscured height  $H$  and relative humidity  $RH$

Methodology:

1. Rain rate  $R$  aloft is estimated using the  $R(Z, Z_{DR})$  relation
2. Rain rate at the surface is estimated using lookup tables computed for different combinations of  $Z_{DR}$ ,  $RH$ , and  $H$

## Summary of Taiwan tests

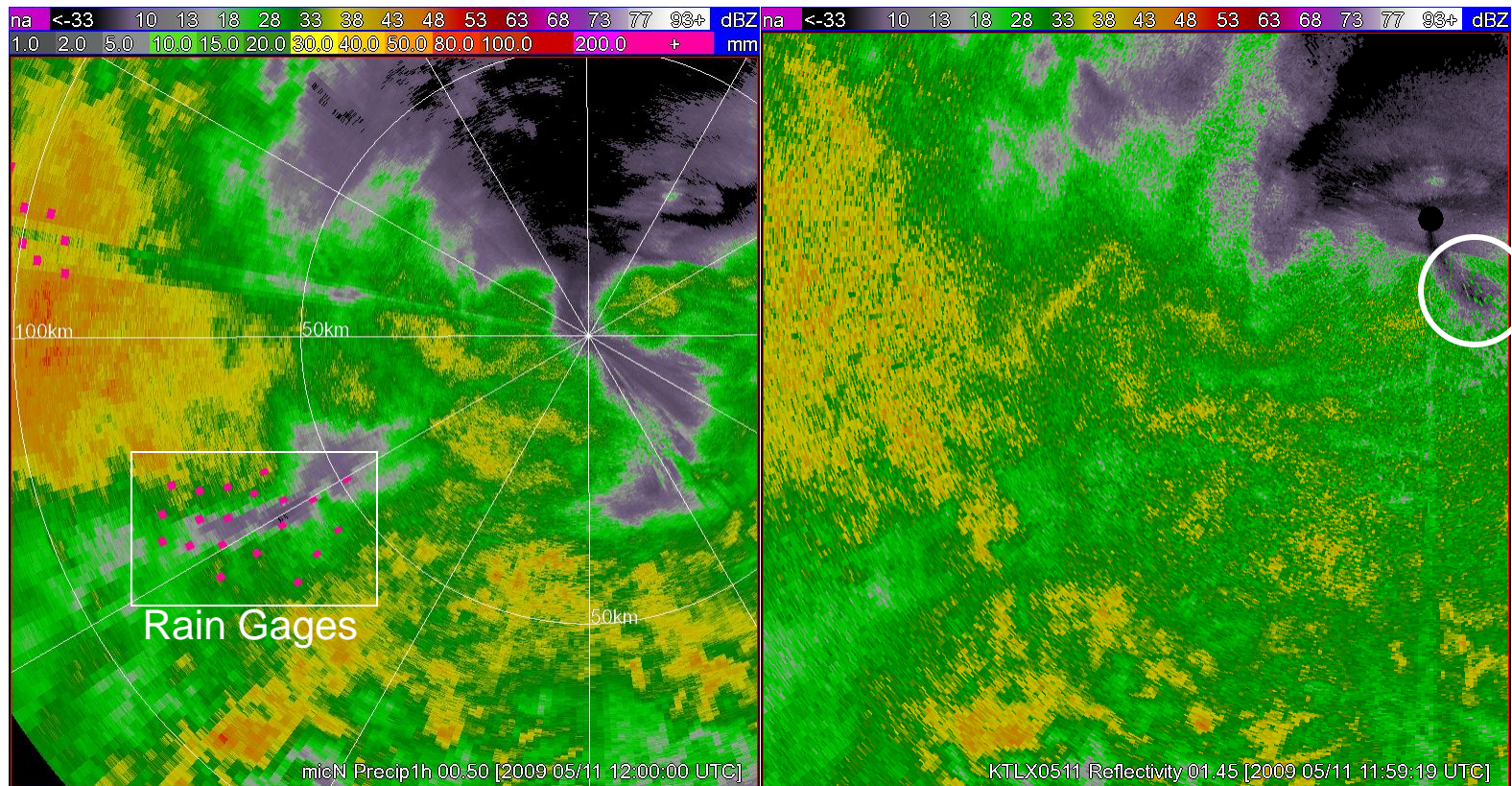
1. Polarimetric algorithms perform better than the conventional  $R(Z)$  (optimized for US) because they are less affected by DSD variability and ground clutter contamination
2. Geometrical correction of partial beam blockage based on DEM leads to incremental improvement
3. The performance of rainfall algorithms (especially of  $R(Z)$ ) is worse at lowest elevation (even in lightly blocked areas) due to contamination from ground clutter
4. Statistically, the  $K_{DP}$  – based algorithm didn't show apparent improvement compared to  $R(Z, Z_{DR})$  at S band for examined rain events

## Suggestions for decision-makers

1. The Taiwan dataset is too small for comprehensive validation of the procedure for rainfall estimation in the presence of beam blockage, hence, no decision on operational implementation of such a procedure can be made at the moment
2. Nevertheless, one of the modules of this procedure, namely, geometrical correction of radar reflectivity factor based on DEM can be recommended for operational implementation
3. More validation studies are required in the complex terrain areas containing dense and well calibrated raingage networks. Possible venues are:
  - utilization of mobile X-band polarimetric radars in Western US (HMT)
  - utilization of operational C-band polarimetric radars and gage network in Taiwan via collaboration between NSSL and TWB

# III. Polarimetric rainfall estimation in the areas affected by ground clutter filtering

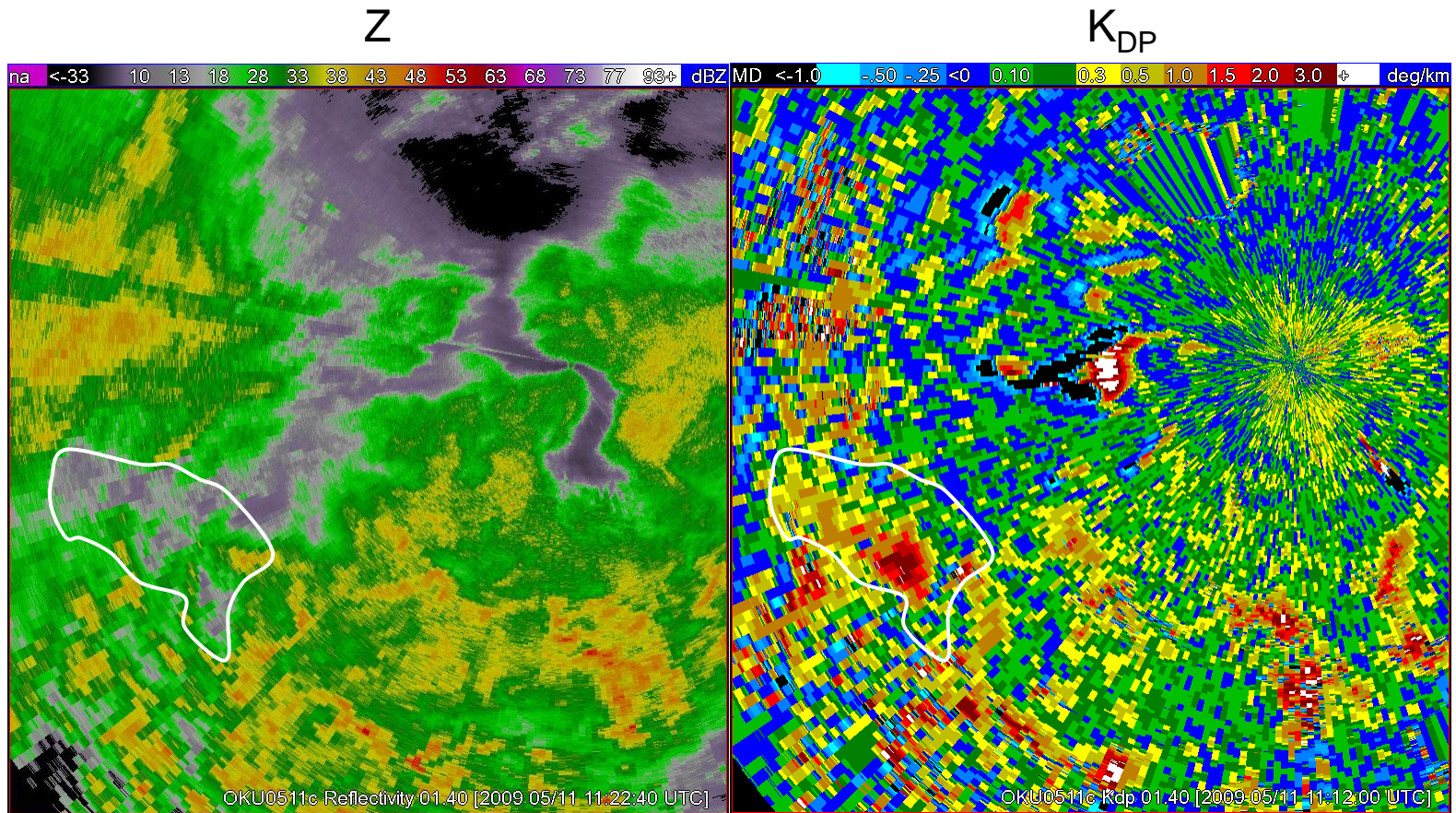
Effect of ground clutter filtering in the OU PRIME field of radar reflectivity



OU'

KTLX

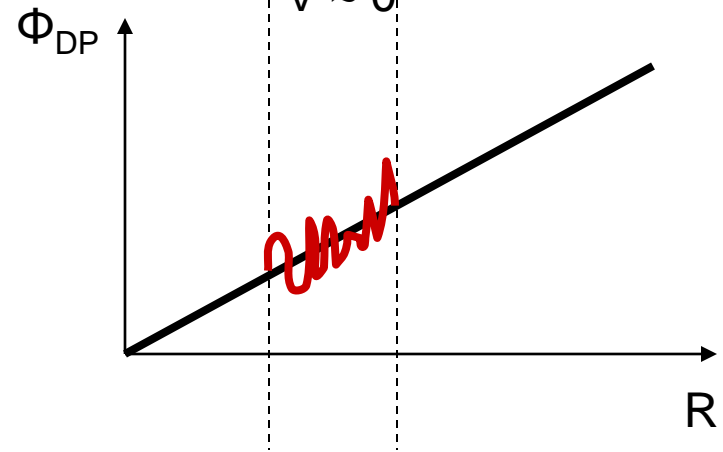
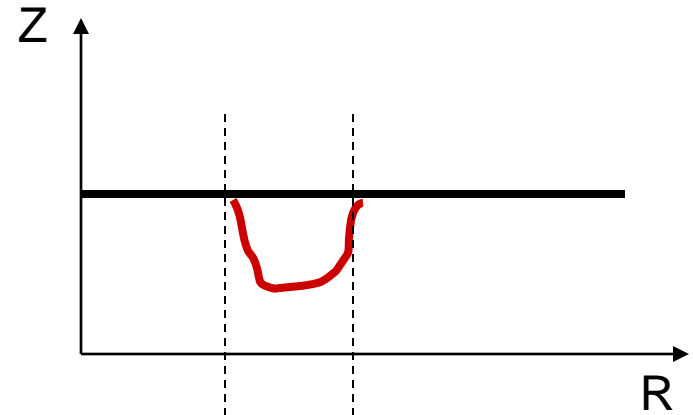
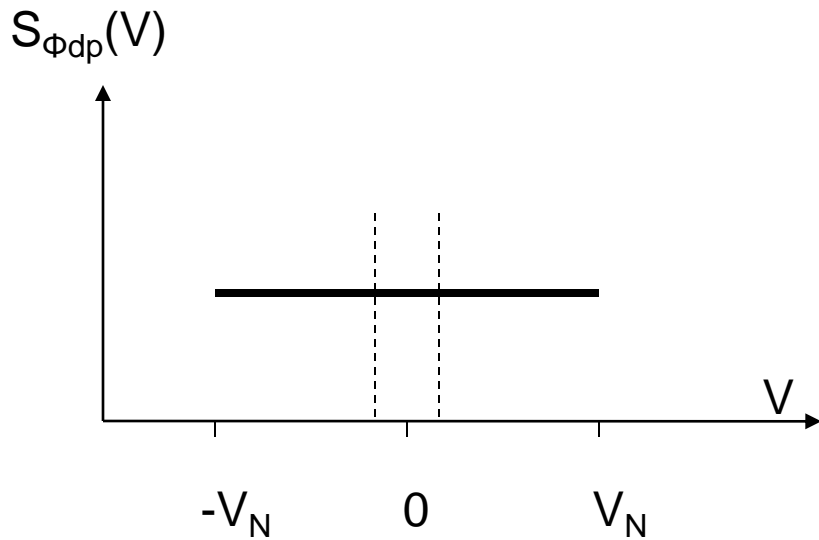
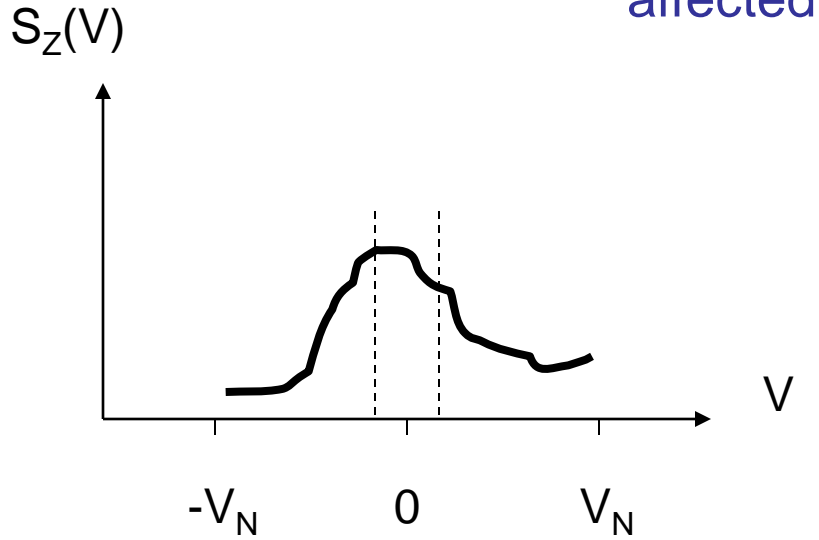
$K_{DP}$  is not affected by ground clutter filtering and exhibits maximum where  $Z$  is reduced due to application of clutter filter at  $V \approx 0$  m/s



Elev. Angle=1.4°



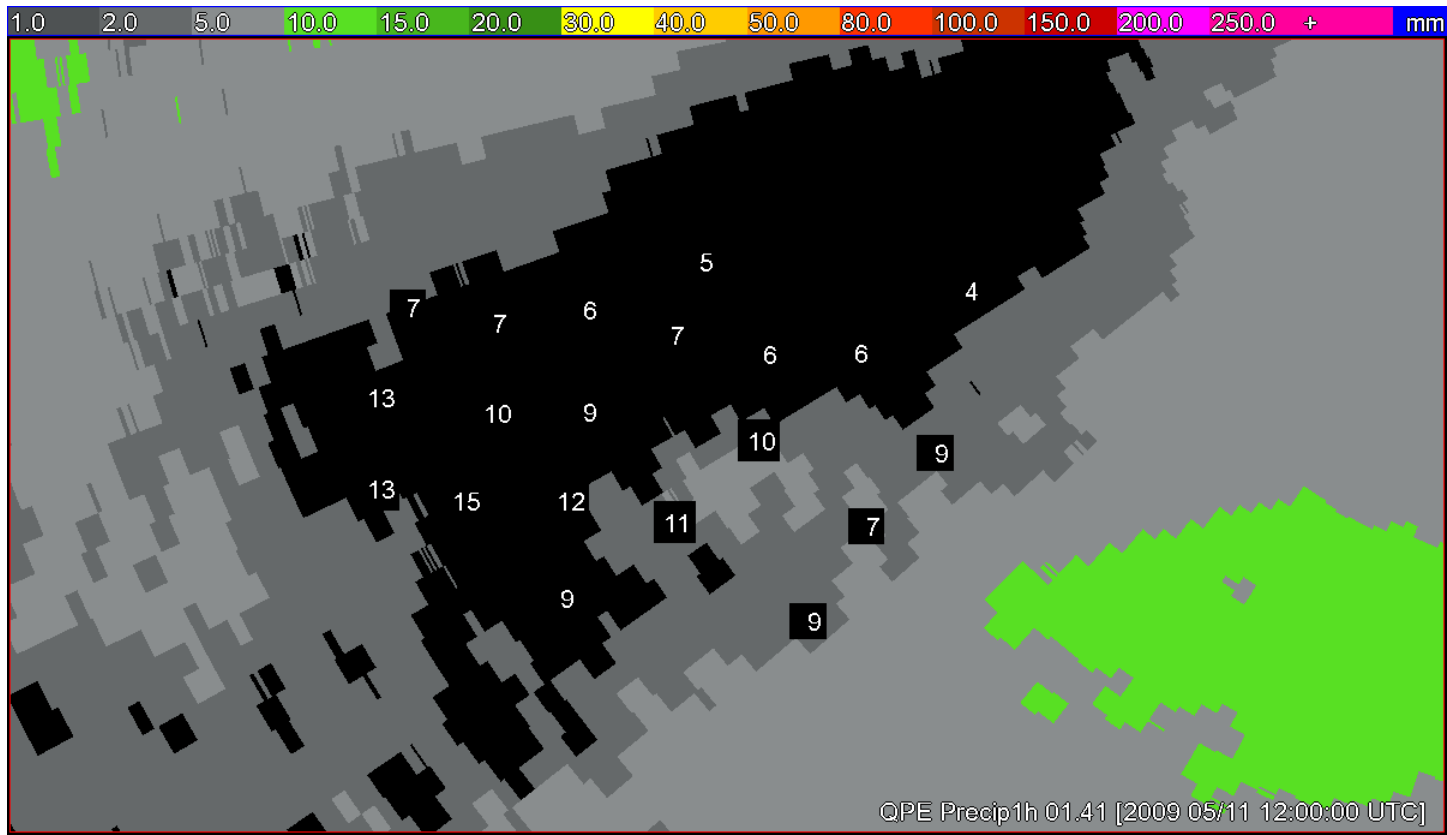
# Radar reflectivity and differential phase are differently affected by notch filter



Z is biased, whereas  $\Phi_{DP}$  is not

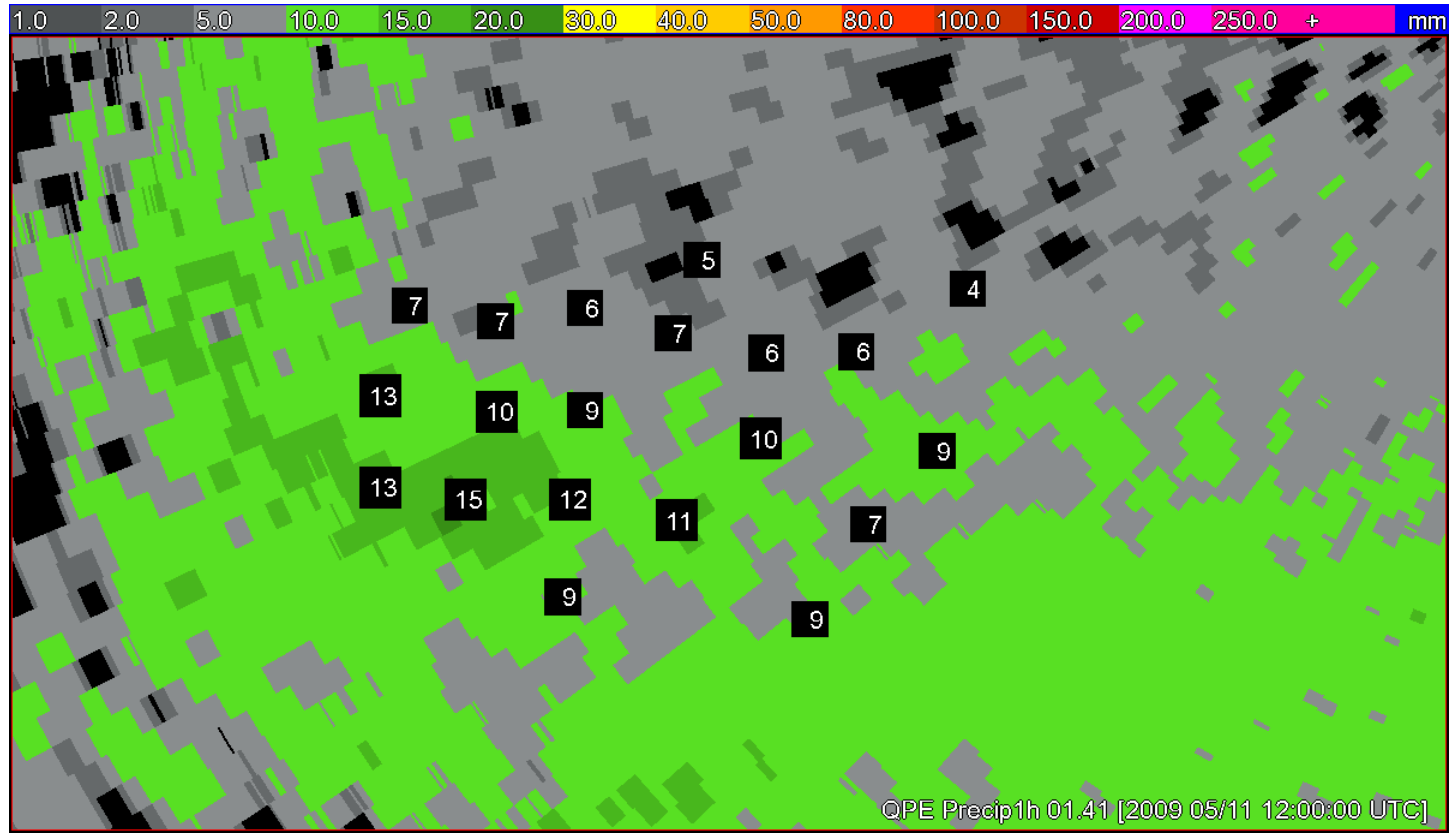
# R(Z) vs rain gages

20090511-12:00 UTC



# R( $K_{DP}$ ) vs rain gages

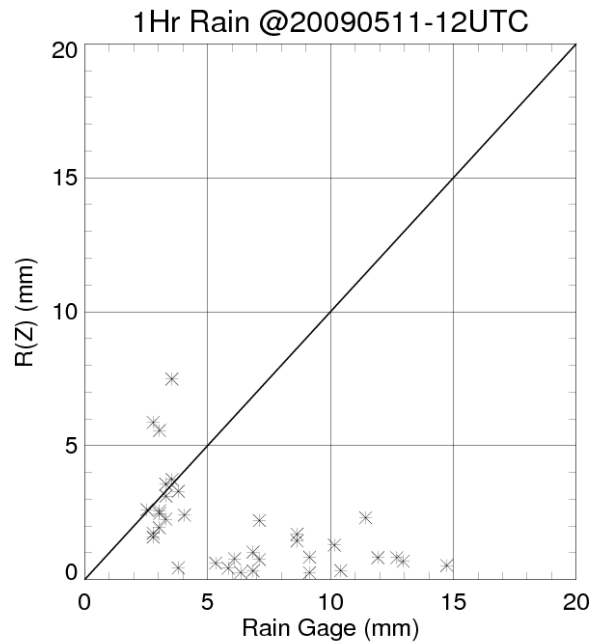
20090511-12:00 UTC



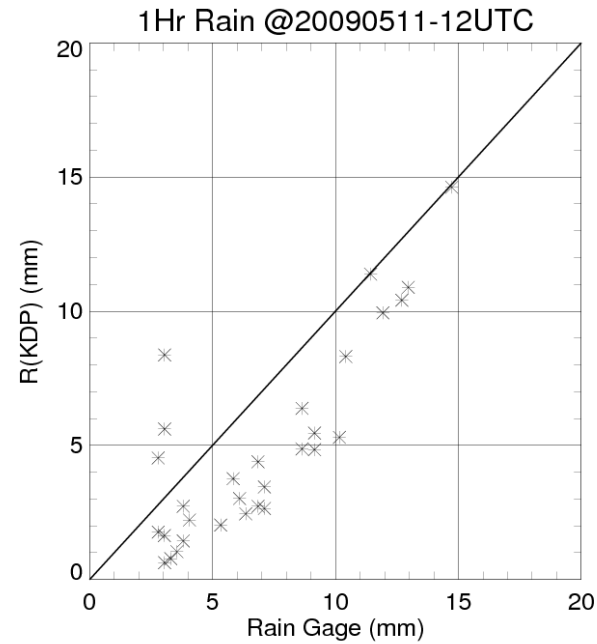
# Hourly radar totals versus gage totals for R(Z) and R(K<sub>DP</sub>)

2009/05 11-12:00 UTC

R(Z)

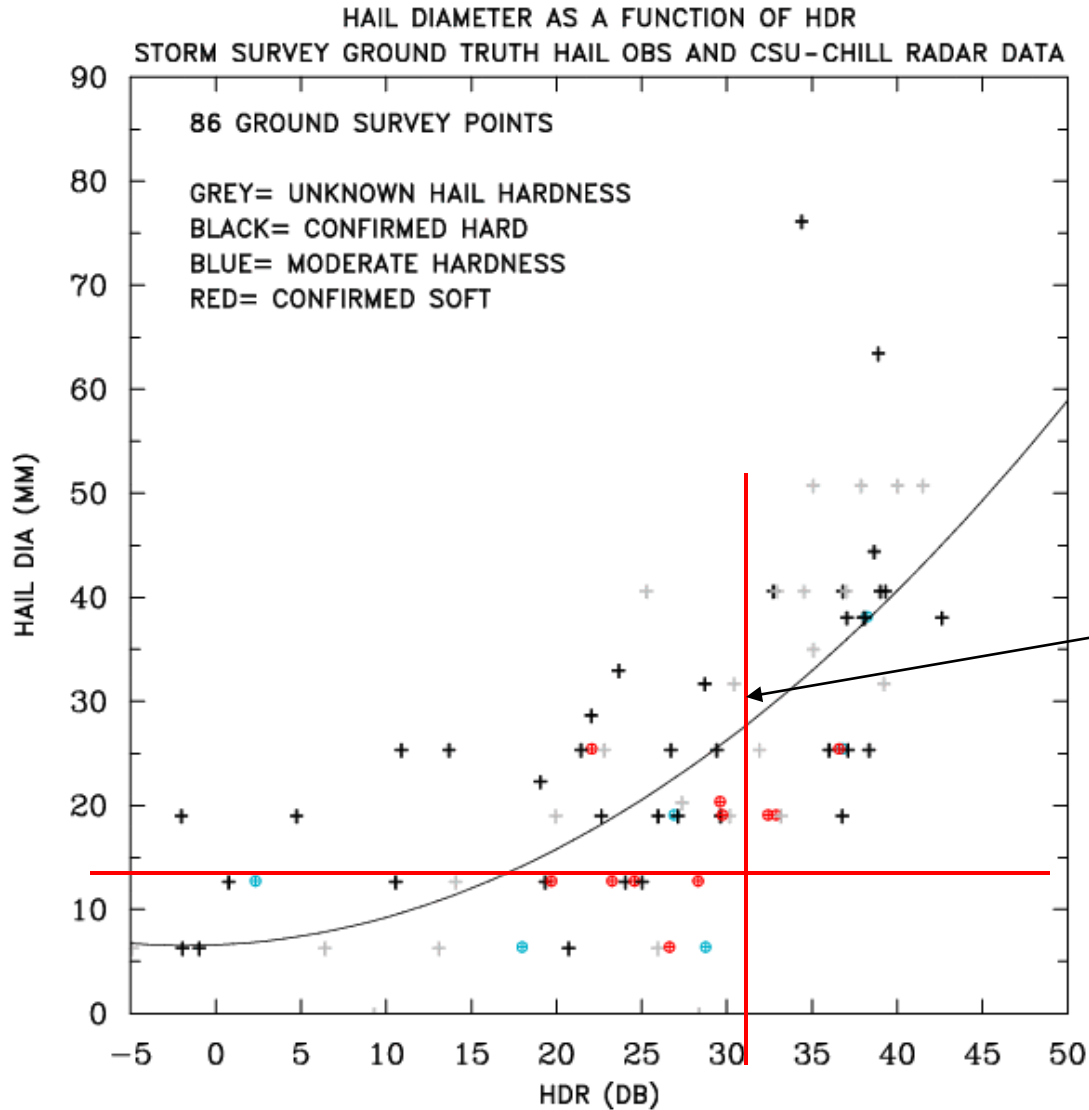


R(K<sub>DP</sub>)



Elev=1.41°

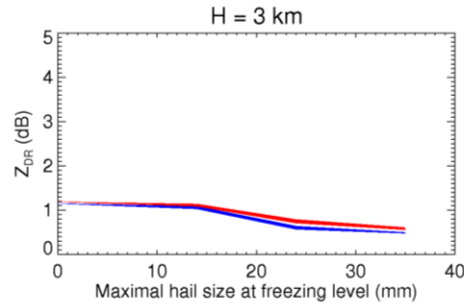
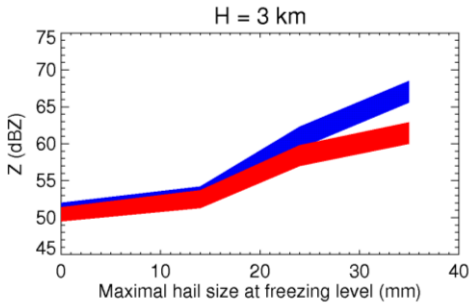
# Modification of HCA to discriminate between small and large hail



$$H_{dr} = Z - f(Z_{dr})$$

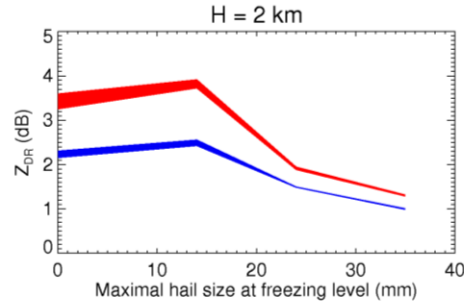
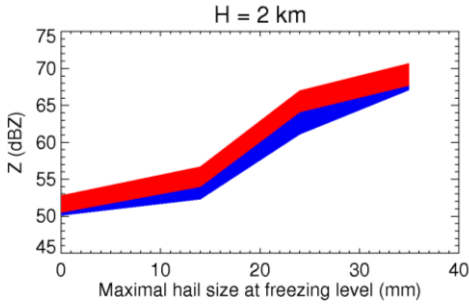
If  $H_{dr} > 30$  dB  
Hail size  $> 20$  mm  
this gave best  
detection of verified  
structural damage

# Dependencies of $Z$ and $Z_{DR}$ on maximal hail size at the freezing level for various parameters of size distribution of ice aloft



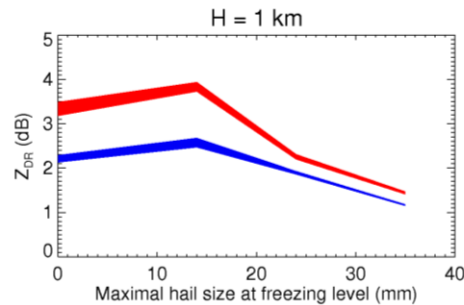
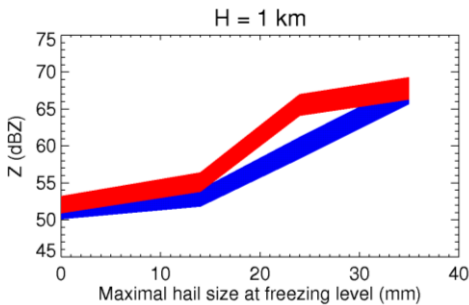
— S band  
— C band

These dependencies are strong functions of height

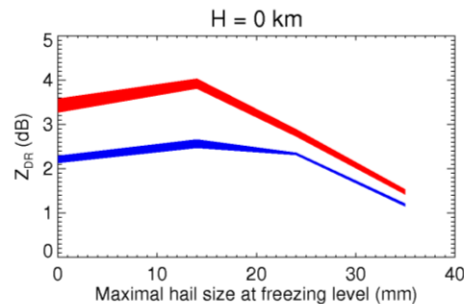
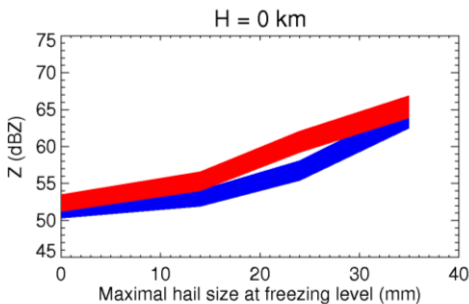


$Z(C)$  can be higher or lower than  $Z(S)$  depending on hail size and height

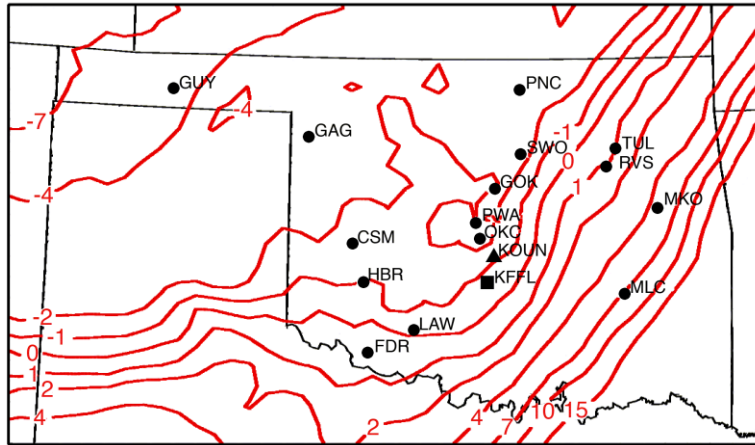
$$Z_{DR}(C) > Z_{DR}(S)$$



The presence of smaller hail tends to increase  $Z_{DR}$  at both radar wavelengths

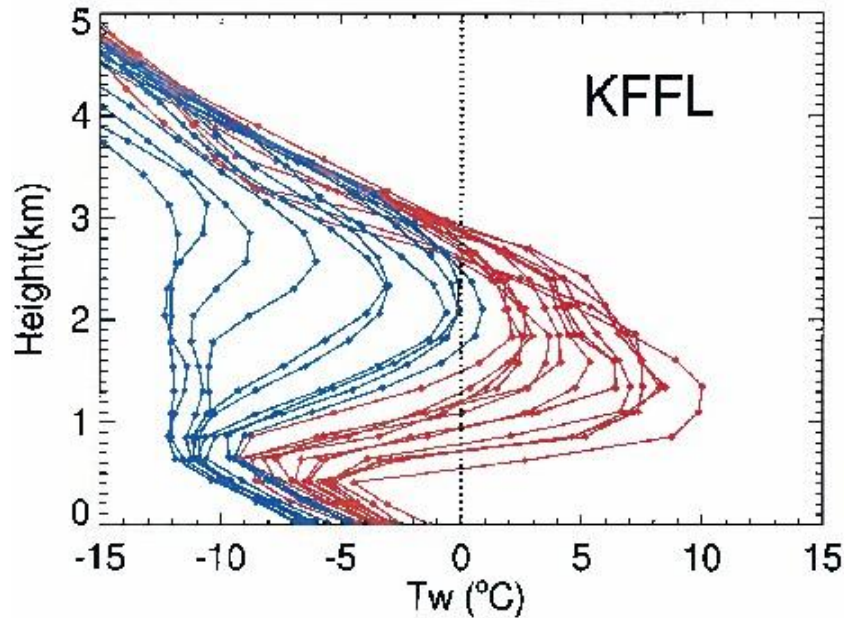


# Developing HCA for transitional winter weather



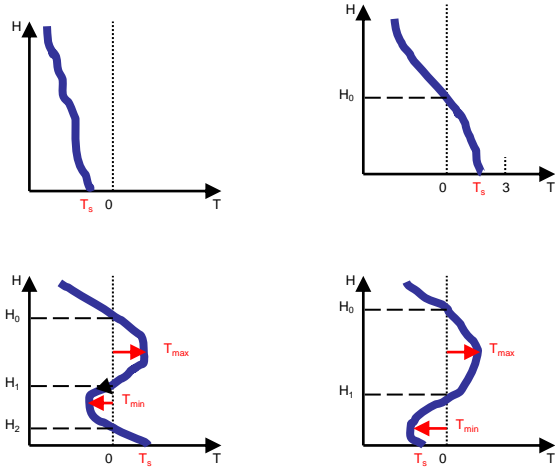
11/30/2006

Surface isotherms  
during passage of cold  
front



Evolution of the  
vertical profile of wet  
bulb temperature

# “Background” hydrometeor classification using vertical profile of wet bulb temperature



## List of classes

HR – heavy rain

RA – light / moderate rain

WS – wet snow

DS – dry snow

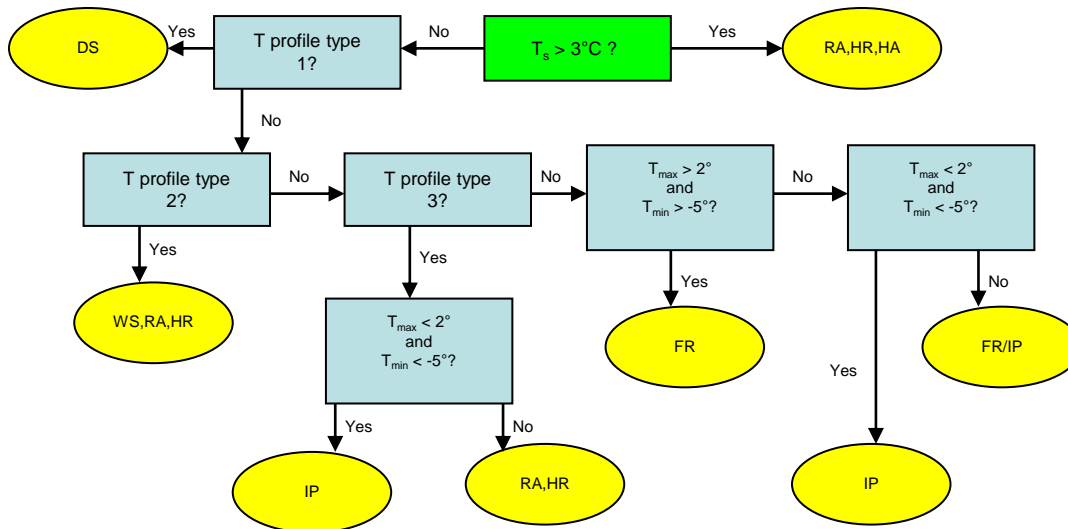
CR – crystals

FR – freezing rain

FR/IP – freezing rain / ice pellets

IP – ice pellets

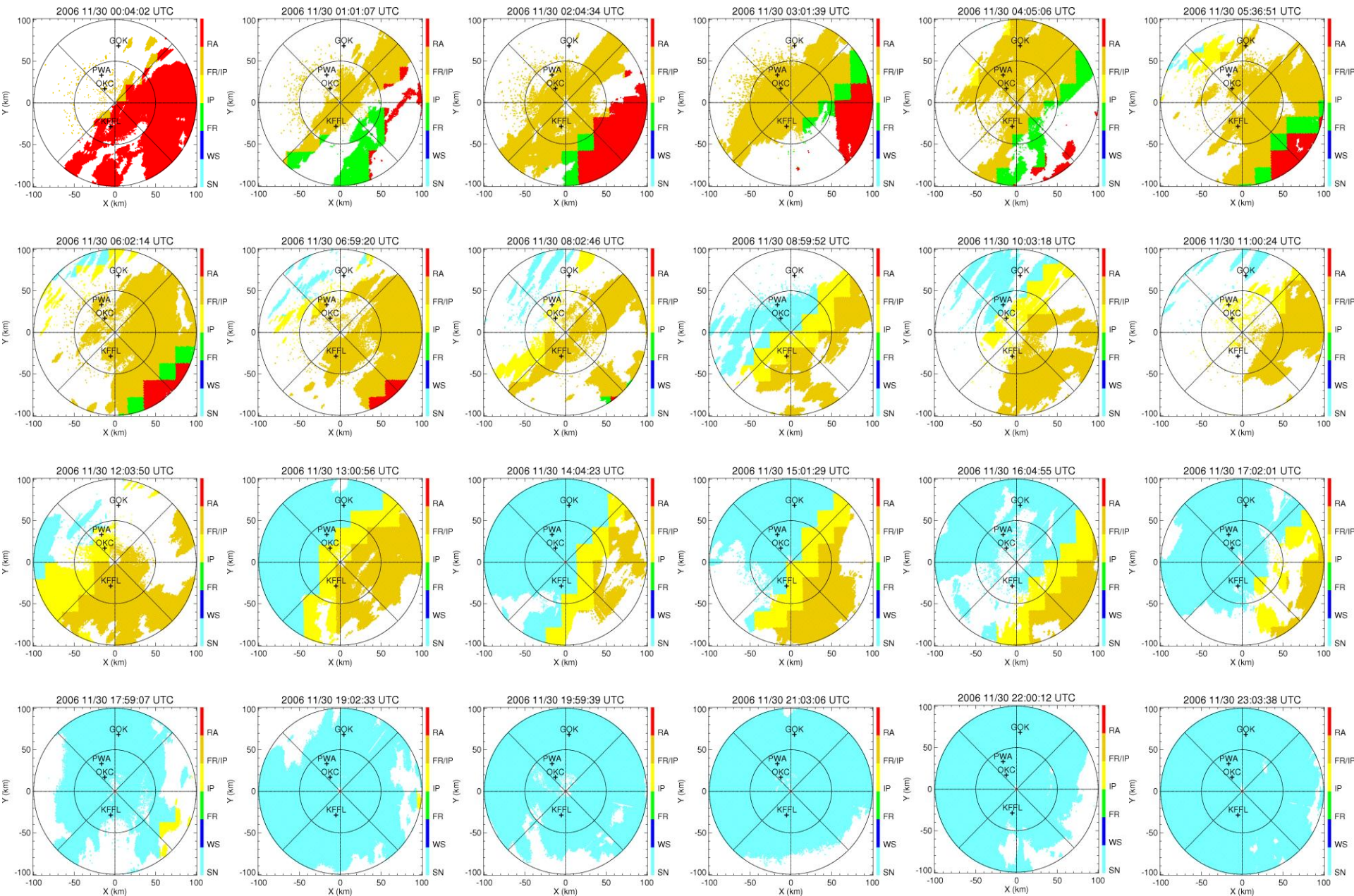
## Four types of vertical profiles of wet bulb temperature



Logistic for determination of precipitation types

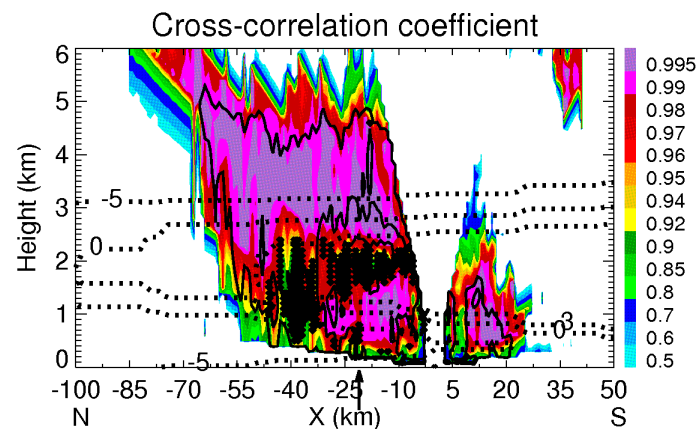
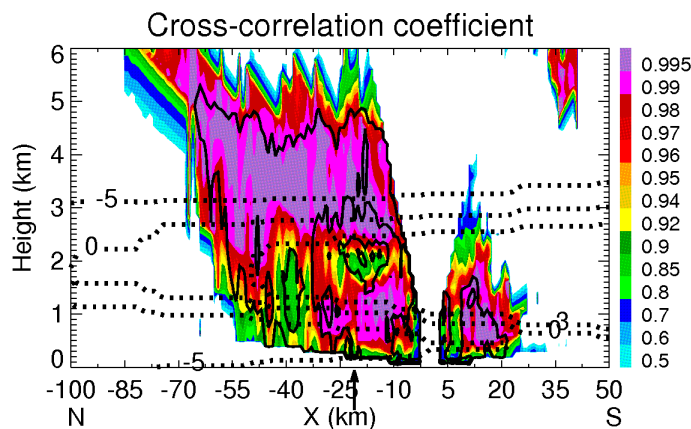
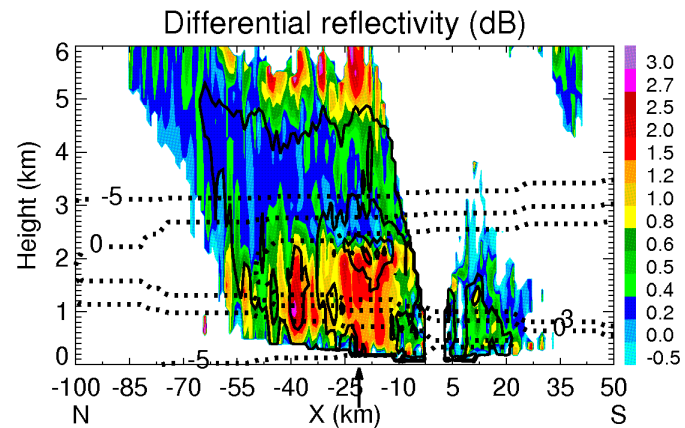
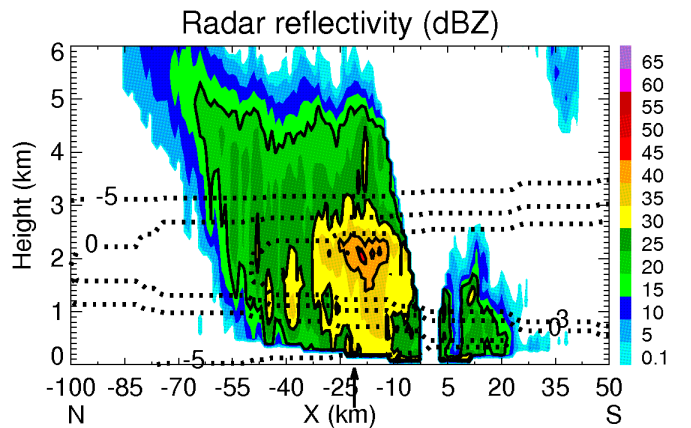


# Background classification



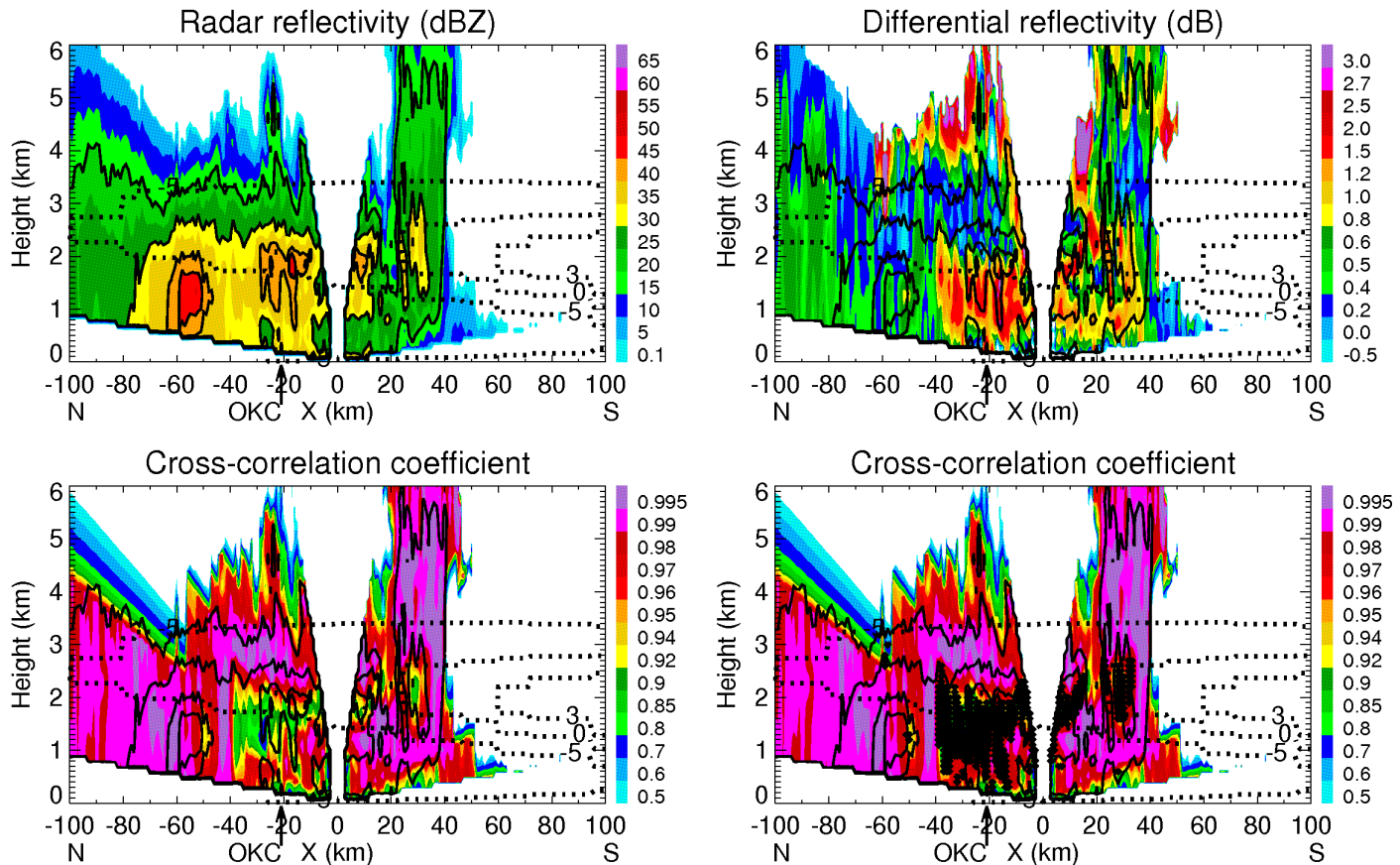
# Example 1

## Radar and RUC model outputs are consistent

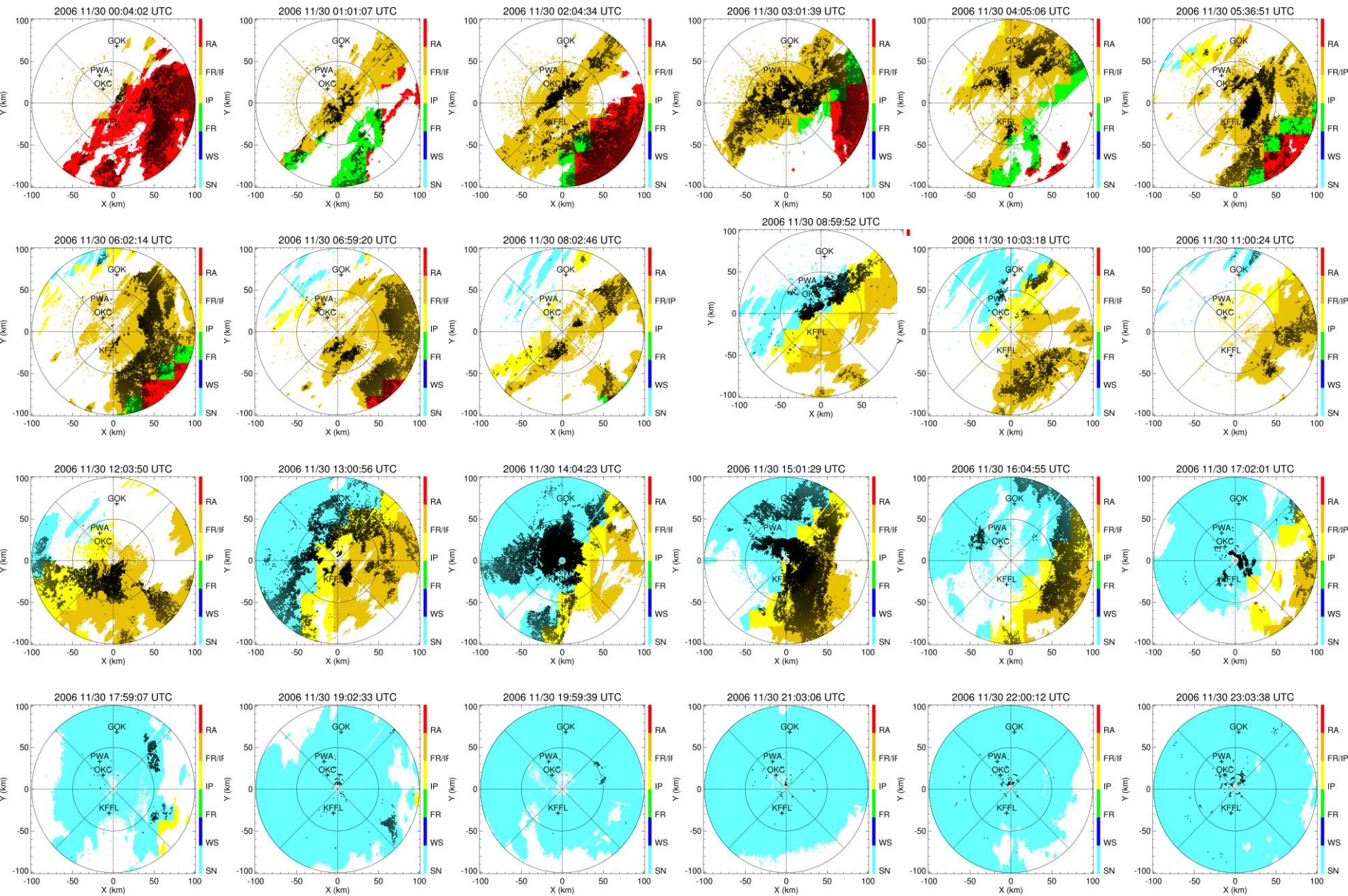


## Example 2

### Radar and RUC model outputs are inconsistent



# Background classification with radar signatures of the melting layer overlaid



# Ground validation of winter HCA

